APPENDIX T

TM: Owens Lake Isotope Study (September 2012)

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September 2012

Owens Lake Isotope Study Technical Memorandum







Owens Lake Groundwater Evaluation Project Isotope Study Technical Memorandum

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LIST OF ACRONYMS AND ABBEVIATIONS

¹² C	carbon-12
¹³ C	carbon-13
¹⁴ C	radiocarbon
¹⁶ O	oxygen-16
¹⁸ O	oxygen-18
³ Н	tritium
³ He	a naturally-occurring rare isotope of helium
⁴He	a naturally-occurring abundant isotope of helium
Ar	argon
CI	chloride
CSD	community service district
D	deuterium
DIC	dissolved inorganic carbon
fmsl	feet above mean sea level
GAMA	Groundwater Ambient Monitoring and Assessment
GMWL	Global Meteoric Water Line
Н	hydrogen
Не	helium
Kr	krypton
LADWP	Los Angeles Department of Water and Power
meq	milli-equivalents
mg/L	milligram per liter
NA	sodium
Ne	neon
NG	Noble Gas
OLGEP	Owens Lake Groundwater Evaluation Project
PDB	PeeDee Belemnite
per mil	parts per thousand
pmc	percent modern carbon
R	isotope ratio
SAP	Sampling and Analysis Plan
SMOW	Standard Mean Ocean Water
STP	standard temperature (20°C) and pressure (1 atmosphere)
ТМ	Technical Memorandum
TU	tritium units or number of ³ H atoms per 10 ¹⁸ hydrogen atoms
uS	microSiemens
USGS	U.S. Geological Survey
Xe	xenon
δ ¹³ C	carbon-13 as a per mil deviation from standard
δ ¹⁸ Ο	oxygen-18 as a per mil deviation from standard
δD	deuterium as a per mil deviation from standard

EXECUTIVE SUMMARY

The purpose of this Technical Memorandum (TM) is to present the results of the Isotope Study conducted under Task 401.1.9.2 of the Owens Lake Groundwater Evaluation Project (OLGEP). The location for this investigation is the Owens Lake study area bounded by approximately Lone Pine to the north and Haiwee Reservoir to the south. Groundwater, surface water, and/or spring water from twenty-eight sites was sampled in July 2012 for a combination of:

- Cations and anions,
- Stable isotopes,
- Radiocarbon,
- Tritium, and/or
- Noble gas.

Isotope measurements in groundwater are conducted for two specific reasons: 1) to determine the source region of groundwater recharge and 2) to determine the age of the water. Prior to sampling, a Sampling & Analysis Plan was developed (MWH, 2012) that outlined the specific goals of the study, including:

- 1) Determine how uniform groundwater isotopic signatures were in groundwater beneath Owens Lake,
- 2) Define the isotopic character of springs and shallow wells on the west side of Owens Lake and how they may be related to deeper groundwater beneath the lake, and
- 3) Determine the isotopic character of shallow groundwater on the east side of the lake.

The SAP provided an initial interpretive model based on pre-existing data in the study area. Based on the results of the sampling and analysis program, this TM articulates findings relative to the goals of the study and further advances the interpretation of recharge sources and ages and how they inter-relate.

To summarize, this limited isotopic study was successful in meeting the goals outlined in the SAP and improved the hydrogeologic conceptual model in several ways. Selected key findings that contribute to an update of the conceptual hydrogeologic understanding of the study area include:

- The δD-δ¹⁸O results demonstrate that the dilute west side waters are recharged by stream channel infiltration originating from the higher elevations of the eastern Sierra Nevada. This is further supported by the similarity in isotope abundances between measured creek samples and local groundwater and springs.
- 2) The δD-δ¹⁸O of sampled OLGEP monitoring wells (except site DWP-7) demonstrates that they are not recharged locally, but rather their lower isotopic values are consistent with isotopic values of recharge sources further north. In other words, water in OLGEP monitoring wells is consistent with a down-valley flow source.

- 3) Based on the δD-δ¹⁸O results, it can be concluded that groundwater at OLGEP well site DWP-6 located in the southeastern portion of Owens Lake was recharged by significantly evaporated surface water. This suggests that historically, ponded water may have recharged groundwater aquifers locally in the southern portion of the Owens Lake area.
- 4) Based on the geographic distribution of groundwater δ^{18} O values in the OLGEP study area, mapped faults beneath the western side of Owens Lake are likely limiting groundwater movement between west side recharge and down-valley flow sources.
- 5) The δD-δ¹⁸O values for springs and shallow wells located on the east side of Owens Lake are generally the same as those for OLGEP monitoring wells recharged by downvalley flow, which indicate a similar isotopic and geochemical character. However, there are some differences observed in the isotopic values, such as in DWP-3/T899, that suggest locally-recharged groundwater does contribute to some recharge on the east side.
- 6) Groundwater ages measured for springs and shallow wells on the west side of Owens Lake were found to be less than 300 years old.
- 7) Groundwater from three sites (Cottonwood Spring NW, LADWP Cottonwood Plant Well, and Cartago-CSD) was found to be less than 50 years old.
- 8) Calculated ¹⁴C groundwater ages ranged from approximately 6,800 years to greater than 40,000 years in OLGEP monitoring wells that are recharged by down-valley flow, as well as sample sites on the east side of Owens Lake.
- The ¹⁴C measured in OLGEP well site DWP-9 was heavily influenced by active methane production at depth, and age-dating is highly uncertain.

Future isotopic and general water quality sampling and analysis may prove useful for shedding light in key areas:

- Previous sampling by GBUAPCD, LADWP, and others utilized varying sample suites for cations and anions. A uniform sampling and analysis plan for general minerals (cations and anions) across the study area could provide a useful baseline of water quality. General minerals could also be used as a simple monitoring tool for potential migration of saline water into less saline groundwater following a change in groundwater use patterns.
- 2) Better resolution for recharge sources of down-valley could be provided by additional measurements of δD and $\delta^{18}O$ conducted on a number of surface and groundwater sites in suspected recharge areas and along flow-paths leading to the OLGEP study area.
- 3) Better age-dating of down-valley flow may be provided by sampling for ¹⁴C and noble gases at OLGEP monitoring well sites DWP-1 and DWP-11 (or similarly accessible wells with well-known depths) and/or down-valley flow sites east of the Owens River where there is well construction data and control on the depth of sampling. However, the sampled wells should be absent of free gas of the type that was observed in site DWP-9.
- 4) Faulting in the eastern side of Owens Lake that influences groundwater migration could be further defined by additional samples collected for δD and $\delta^{18}O$, particularly for any wells that might be upgradient of the lake level or adjacent wells with distinctly

different depths. This will help establish the influence of any locally recharged groundwater flowing towards Owens Lake, and whether this locally-recharged groundwater comingles with down-valley flow.

5) Because $\delta D - \delta^{18}O$ measurements are diverse enough among the different recharge sources, additional measurements could be used as a simple and inexpensive groundwater recharge mapping tool.

1.0 Introduction and Study Goals

The Owens Lake Groundwater Evaluation Project (OLGEP) was initiated in March of 2009. This project involves 10 tasks:

Task 401.1.1	Data Compilation
Task 401.1.2	Data Evaluation and Identification of Data Gaps
Task 401.1.3	Assist in the Collection of Field Data
Task 401.1.4	Update Conceptual Hydrogeologic Model
Task 401.1.5 and 401.1.10	Numerical Groundwater Model Update and Development
Task 401.1.6	Model Simulations and Alternative Analysis
Task 401.1.7	Develop and Implement a Public Outreach Plan
Task 401.1.8	Project Meetings/Final Report
Task 401.1.9	Geophysical Analysis and Isotope Analysis

The first six tasks are generally sequential, while the last two tasks have periodic activity throughout the project. A preliminary groundwater model and associated model documentation technical memorandum (TM) was completed in February 2012 for the OLGEP study area (see **Figure 1**) and for review by Partner Agencies and the Blue Ribbon Panel. The model was used to evaluate alternatives for production of groundwater at Owens Lake. At present, this model is being updated and will be used for defining and selecting a preferred alternative.

Recent meetings with Partner Agencies and other stakeholders highlighted the need to characterize groundwater recharge sources and age. This is particularly important in understanding the source waters for springs and seeps, and ultimately, improvement of the conceptual understanding and numerical modeling of the effects that pumping has on groundwater discharge areas.

Sampling of surface and groundwater for a variety of constituents (including isotopes) has been used successfully to evaluate the origin of water flowing to springs or seeps, or to characterize mixing (or lack of mixing) of separate hydrostratigraphic zones in other locations in the Owens Valley (MWH, 2004). The installation of 28 new deep monitoring wells during earlier phases of the OLGEP project (hereby referred to OLGEP wells) provides an opportunity to characterize deeper groundwater and correlate this groundwater with surface discharge sources.



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The purpose of Task 401.1.9.2 entitled, "Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas" was to:

- 1) Review previous isotopic and contemporaneous associated general mineral sampling in the OLGEP study area and develop an isotope conceptual model for the study area.
- 2) Based on this data, develop a sampling and analysis plan (SAP) (MWH, 2012) to characterize the source waters of the most significant aquifers at depth and the most significant groundwater discharge locations, better define recharge amounts and sources, and to understand the source water or origin of springs.
- 3) Perform targeted sampling of selected monitoring wells and surface discharge areas at locations shown on **Figure 1**.
- 4) Evaluate and report the sampling results in a technical memorandum (this document) detailing new information provided by the sampling that improves the understanding of the age and origin of deep groundwater and selected springs and seeps.

Implementation of this task serves as a pilot study to demonstrate the feasibility of isotopes for characterization of spring sourcing in the OLGEP study area and is not intended to be an exhaustive investigation.

This TM represents the final deliverable of the Isotope Study (Item No. 4 above). The TM takes into consideration and is complementary to previous isotopic and general chemistry sampling performed by Lopes (1988), Font (1995), MWH (2004), Great Basin Unified Air Pollution Control District (2009), Densmore et al. (2009), and LADWP (reported by MWH, 2011).

2.0 Background

Isotope measurements in groundwater are conducted for two specific reasons:

- To determine the source region of groundwater recharge, and
- To determine the age of the water.

The most common isotope measurements for determining recharge source are deuterium (D) and oxygen-18 (¹⁸O). These measurements prove useful because they are intrinsic properties of the water molecules; they are stable and do not undergo atomic change as in radioactive isotopes; and their abundances systematically vary in atmospheric precipitation with changing latitude (i.e., temperature) and elevation.

For age-dating of groundwater, the most commonly used tools are radioactive isotopes of tritium (³H) and radiocarbon (¹⁴C). Tritium has a radioactive half-life of 12.43 years, making it ideal for dating recharge less than 50 years old. Radiocarbon has a half-life of 5,730 years and has been used to distinguish much older groundwater at depth, particularly in regional groundwater systems. Inherent problems with ³H age dating is the fact that accurate ages require knowing how much ³H was in the water at the time of recharge. Unfortunately, with surface testing of nuclear weapons in the 20th century, atmospheric ³H varied over an order of magnitude, heavily influencing input values of recharging groundwater. Consequently, the presence or absence of ³H in a groundwater has been the default method for determining whether recent recharge (<50 years old) is incorporated in particular water, but with no particular date of recharge.

Radiocarbon is a common tool in archeological and tree ring studies, and has been used extensively in groundwater. The ¹⁴C is measured in the dissolved inorganic carbon (DIC). Unfortunately, DIC can easily undergo reactions in aquifers such as precipitating carbonate minerals, dissolving ¹⁴C-absent carbonate minerals within the host lithology, or simply exchanging with them. When carbonate is present in soils or aquifer rocks (very common in the western US), DIC will react readily with them, resulting in decrease of the ¹⁴C abundance that is unrelated to radioactive decay. The stable isotope carbon-13 (¹³C) can be used to determine the extent of these reactions, providing they are isotopically distinct from the original DIC ¹³C abundance at the time of recharge. Note that for ³H, geochemical reaction is not a problem because ³H is part of the water molecule (see Clark and Fritz, 1997 for additional background information).

More recently, the measurement of noble gases (helium, neon, krypton, and xenon) that naturally dissolve in recharging groundwater have provided additional recharge source and agedating methods. An accurate measurement of helium isotopes allows the determination of the amount of ³H that has already decayed because ³H radioactively decays to the stable helium-3 (³He) isotope. Consequently, the amount of ³H that was present at the time of recharge can be calculated and accurate age dates can be determined. Furthermore, the amount of stable helium-4 isotope (⁴He) dissolved in groundwater can be a semi-guantitative age-dating method of a similar range as ¹⁴C. The ⁴He originates from steady-state accumulation in groundwater due to production and diffusion upward of uranium and thorium alpha particle decay (alpha particles are ⁴He atoms) in crustal rocks. Note that in areas of active volcanoes, the ³He and ⁴He can be significantly altered by emission of volcanic gases in the subsurface. This will usually dissolve an abnormal amount of He in the groundwater with an excess of ³He compared to crustal rocks. Lastly, because the absolute noble gas abundance dissolved in water is temperature dependent, a method has been developed to also determine the recharge temperature of a groundwater. In the case of geographic areas like Owens Valley, that recharge temperature will be largely elevation dependent (see Mazor, 2003 for additional background information).

An accepted approach for reporting the abundance of D, ¹⁸O, and ¹³C entails reporting them as a ratio to their abundant isotope in the element (H, ¹⁶O, and ¹²C) and converting the ratio to a parts per thousand (per mil) deviation from internationally recognized standards. This per mil deviation defined as

$$\delta = \left(\frac{R}{R_s} - 1\right) 1000,$$

where δ (del) is the per mil deviation, R is the isotope ratio (D/H, ¹⁸O/¹⁶O, or ¹³C/¹²C) measured in a sample, and R_s is the isotope ratio of the standard. For D and ¹⁸O, the standard is Standard Mean Ocean Water (SMOW), and for ¹³C is PeeDee Belemnite (PDB, a marine calcium carbonate). Because continental water is usually depleted in D and ¹⁸O relative to ocean water, the δ values are commonly negative. The remainder of this TM will refer to stable isotope results in δ form and per mil as commonly reported in the literature.

3.0 Previous Work

Table 1 lists the reports reviewed as part of this study. Each report is briefly summarized herein and was used to develop a preliminary interpretive model presented in the SAP (MWH, 2012) and the updated interpretive model presented in this TM.

Table 1
Summary of Previous Isotopic Studies Reviewed

Study	Isotope Measurements	Basis
Inyo/LA Geochemical Cooperative Study (MWH, 2004; Bassett et al., 2008)	δD, δ ¹⁸ O, δ ¹¹ B, δ ³⁴ S, 3H	Investigated possible influence of groundwater pumping on associated spring flows and provide supportive evidence for hydrogeological model.
U.S. Geological Survey (Densmore et al., 2009)	δD, δ ¹⁸ O, 3H, δ ¹³ C, ¹⁴ C, NG	Performed for State Water Resources Control Board's GAMA program: http://www.swrcb.ca.gov/gama/
Font (1995)	δD, δ ¹⁸ O, δ ¹³ C, ¹⁴ C, ⁸⁷ Sr/ ⁸⁶ Sr	Trace sources of recharge and develop evaporation model for Owens Lake.
Lopes (1988)	δD, δ ¹⁸ Ο	Develop hydrologic mass balance of Owens Lake and process of salt crust formation.
Ingraham and Taylor (1991)	δD, δ ¹⁸ Ο	Sampled precipitation along transects from coastal California to western Nevada.
Friedman et al. (1970; 1992; 2002)	$\delta D, \delta^{18} O$	Comprehensive precipitation collection and measurements through southeastern California and Nevada.

The most comprehensive isotope investigation of groundwater in the Owens Valley was conducted as part of the Inyo/LA Geochemical Cooperative Study, which investigated the possible influence of groundwater pumping on associated spring flows and provided additional supportive evidence for hydrogeological modeling. The study was confined to the upper and middle parts of the Owens Valley and does not overlap the lower Owens Valley/Lake area in this present study. However, some important conclusions were developed in this study. For instance, the presence of measureable ³H in many of the well waters collected indicated that relatively young groundwater was being produced. Many wells had large screened intervals that intersected shallow depths, supporting the conclusion that sampled groundwaters were a mixture of younger and older water. The presence and absence of ³H in springs was diagnostic of deeper versus shallow groundwater discharge. However, the δD and $\delta^{18}O$ values were not diagnostic of different recharge sources, nor based on the report's conclusion, indicative of evaporation. Reliance on the $\delta^{11}B$ and $\delta^{34}S$ values subsequently provided much of the geochemical evidence of similarities and differences between spring discharge and wells, as well as the influence of local faults on groundwater flow.

A more recent groundwater investigation by the U.S. Geological Survey (USGS) collected groundwater from 53 individual wells throughout the Owens Valley, including the Owens Lake area (Densmore et al., 2009). This work was performed under the State Water Resources Control Board's Groundwater Ambient Monitoring and Assessment program (GAMA). Analysis included a full suite of isotope data involving δD , $\delta^{18}O$, $\delta^{13}C$, ¹⁴C, and noble gases. All data except for the noble gas results have been published, but data interpretation by the USGS is forthcoming. Nevertheless, data have been tabulated and independently interpreted for the purposes of this study.

Font (1995) and Lopes (1988) independently conducted graduate thesis research of shallow groundwater below Owens Lake. Both studies sought to understand water mass balance, and in the case of Font (1995) to determine source regions of recharge. Font used a method to calculate the extent of water evaporation from the evaporated values of δD and $\delta^{18}O$. Font also used strontium isotope measurements to delineate recharge derived from the Sierra Nevada to the west from that derived from the east. In the case of Lopes (1988), limited stable isotopes of water were measured, and only general statements about possible origins of recharge were proposed. The main body of work was dedicated to evaporative enrichment and salt chemistry.

Several studies measuring δD and $\delta^{18}O$ of atmospheric precipitation in the Sierra Nevada, southeastern California, and Nevada have been conducted. Friedman et al. (1970) did initial work characterizing the change in rain and snow isotope abundances on the western slope of the Sierra Nevada and showed the systematic depletion of these abundances with increasing elevation. Ingraham and Taylor (1991) mapped similar trends for three transects extending from coastal California through Nevada. The southernmost transect passes through the Owens Valley. Friedman et al. (1992; 2002) collected and measured δD and $\delta^{18}O$ values for integrated precipitation samples throughout southeastern California and Nevada over multiple years. They showed that a systematic depletion of values with increasing elevation did not occur systematically in these regions as was observed on the western slope of the Sierra Nevada Mountains.

The preliminary interpretive conceptual model presented in the SAP (MWH, 2012) has been updated using new data and results from this study (Section 7 - Update of Interpretive Model).

4.0 Sample Collection

Samples collected for this investigation involved surface water runoff, springs, production wells, and monitoring wells as summarized in **Table 2** and shown in **Figure 1**. Details of each collection site are included in **Appendix A**, whereby a site summary sheet presents key information by site (location coordinates, elevation, well information, field data recorded, samples submitted to laboratory, air photo location map, and site photos).

Surface runoff was collected at Carroll Creek (De La Cour Ranch) located northwest of Owens Lake and another at Ash Creek further south along the west side of Owens Lake (**Figure 1**). These samples likely represent baseflow conditions for the streams. Samples for δD and $\delta^{18}O$ were collected at these sites. Sampling entailed filling of a 4-oz glass bottle with a conical-shaped plastic insert inside the cap to provide an air-tight seal. Water was collected from flowing currents in the stream bed to avoid bias from pools undergoing evaporation.

Springs and seeps that were sampled are located in all areas surrounding Owens Lake with the exception of the north side. Spring collection was focused on the western side as well as one location to the southeast. Spring sites were chosen to generate isotope data in areas lacking sufficient coverage in previous studies. Four springs in total were sampled. Three of these springs had somewhat of a diffuse discharge, whereas the fourth location had flowing water from an open standpipe. For diffuse sites, the wettest area was determined, usually designated by saturated soil conditions, and a hand-dug hole approximately 3-feet deep was dug. Water from the hole was bailed, refilling rates were observed, followed by sample collection. Samples for δD - $\delta^{18}O$, ³H, and $\delta^{13}C$ and ¹⁴C were collected at these sites. The δD - $\delta^{18}O$ samples were collected in 4-oz glass bottles as discussed above. Tritium was collected in 16-oz glass bottles, and $\delta^{13}C$ and ¹⁴C in a single 8-oz glass bottle, both with the same conical-shaped plastic insert

inside the cap. Samples were filled after completely submerging bottles underwater and filled gently to avoid air entrainment. For water flowing from a stand-pipe, the same collection method was used, with care taken not to entrain air.

Remaining samples were collected from wells. Five of these were domestic supply wells, four of which had pressurized storage tanks. For smaller tanks (i.e., 20-40 gals) the well was operated for a period of time to discharge stored water so that the most recently pumped water was collected. Cartago CSD fed a much larger storage tank that was impractical to discharge, and stored water likely influenced the collected samples. However, all pressurized tanks were isolated from the surrounding air, and exchange with the water was inferred to be minimized. Rio Tinto was also a domestic supply well, but had no pressurized storage tank at the point of sampling. Two wells were dedicated for agricultural purposes and samples were collected from high volume discharge emitted from a large diameter horizontal pipe downstream of the pump.

In addition, samples were collected for general minerals by LADWP on various surface waters, springs, and wells on the west side of the OLGEP. Data from these samples along with previously-generated water quality data in the OLGEP study area were compiled and are briefly discussed in the context of the isotope results.

Monitoring wells that were sampled include new OLGEP monitoring well clusters completed at different depths in the same location. These were drilled by LADWP specifically for the purposes of hydrogeologic characterization and water guality monitoring associated with the OLGEP. All monitoring wells sampled were artesian, and discharge was collected from a sample port through a 90° ball valve assembly. Wells were purged for approximately 20 minutes until field measurement parameters stabilized. In several of the wells, free gas was associated with the discharge. It was concluded that the gas was associated with aquifer water and not related to well construction or water collection. Methane discharge from deep groundwater beneath Owens Lake is not uncommon. Samples were collected by the same method in glass bottles as discussed above. In two of the three wells from each cluster sampled, noble gases were also collected. This was accomplished by connecting to the sampling port a 3/8" barbed taper fitting that was joined to a flexible 3/8" Tygon tubing and sealed with a hose clamp. The other end of the tubing was sealed to a 3/8" copper tube with a hose clamp. The copper tube was positioned into a dual pinch-clamp assembly that can sequentially create a gas-tight closure of the copper by a cold pinch seal. Downstream of the copper was attached an additional 3/8" Tygon tubing connected to a second 90° ball valve. Water was flowed through the tubing and copper tube at a laminar flow rate. The downstream ball valve was closed down far enough to re-dissolve any free gas in the water. All free air from all tubing was removed by agitation followed by pinch clamping the downstream end of the tube. The upstream end was then clamped to seal an air-free sample inside the copper tube between the two clamps. All noble gas samples were collected in duplicate. Noble gas samples were collected only from OLGEP monitoring wells (see Table 2; Figure 1; Appendix A). All samples except DWP-3/T901 were collected successfully with a high degree of certainty that air bubbles were not entrained in the sample. For DWP-3/T901, the free gas present in the water during sampling was excessive, and some level of uncertainty about the guality of the sample was noted.

Table 2Summary of Isotope Study Sampling Locations

Location	Field Label	Spring	Surface	Well	Cations/Anions	Stable I	sotopes	Radiocarbon	Tritium
		opg	Water			δD and $\delta^{18}O$	δ^{13} C		
West Side (from north to south)									
De La Cour Ranch (Carroll Creek)	Carroll Crk		x		1	1			
Northwest Seep	NW Seep	х			1	1	1	1	1
Bartlett Well (BRTLTPW)	Bartlett-CW	х		х		1	1	1	1
Cottonwood Spring NW	Cottonwood Spring	х			1	1	1	1	1
Cottonwood Spring Well	Cottonwood-Well			х		1			
Cottonwood Flume	N/A (Field data only)		x						
LADWP Cottonwood Plant Well	Aqueduct Well			х	1	1	1	1	1
Ash Creek	Ash Crk		x		1	1			
Rio Tinto Well	Rio Tinto-SW			х	1	1	1	1	1
Cartago CSD	Cartago-COM			х	1	1	1	1	1
Deep Wells (from north to south)									
DWP - 9 T896	DWP - 9 T 896			x		1	1	1	1
DWP - 9 T897	DWP - 9 T 897			х		1			
DWP - 9 T898	DWP - 9 T 898			х		1	1	1	1
DWP - 3 T899	DWP - 3 T 899			х		1	1	1	1
DWP - 3 T901	DWP - 3 T 901			х		1	1	1	1
DWP - 6 T912	DWP - 6 T912			х		1	1	1	1
DWP - 6 T911	DWP - 6 T911			х		1	1	1	1
DWP - 6 T913	DWP - 6 T913			х		1	1	1	1
DWP- 7 T908	DWP-7 T908			х		1	1	1	1
DWP- 7 T909	DWP-7 T909			х		1			
DWP -7 T910	DWP -7 T910			x		1	1	1	1
East Side									
Dolomite Well	Dolomite			х		1	1	1	1
AGRP Well	Agr-2	х				1	1	1	1
Keeler CSD Well	KCSD-PW	х		х		1	1	1	1
Horse Pasture Well (Abandoned)	Horse Pasture			x		1	1	1	1
S3 (1) (SE Pieozometer)	S31			x	1	1	1	1	1
Duck 3 Well	Duck-CN			x		1			
Tubman Cement Pond	Tubman-CP	х				1			
				TOTAL:	8	27	20	20	20

Noble Gas	Rationale
	Further characterize water on the west side.
1 1 1 1 1 1 1	Determine the source and age differences between aquifer 1 and aquifer 5, or in other words shallow and deep aquifers.
	Supplemental data for east side: 1) high discharge spring age and source; 2) up gradient well for recharge source determination; 3) groundwater sample furthest east in the basin for comparison with west side
6	

5.0 Analytical Methodology

Samples were packaged and shipped within 48 hours to Isotech Laboratories Inc. for δD , $\delta^{18}O$, $\delta^{13}C$, and ³H analyses, and preparation of samples for ¹⁴C analysis. Noble gas samples were shipped to the Dissolved and Noble Gas Lab in the Geology and Geophysics Department of the University of Utah, Salt Lake City.

The $\delta D - \delta^{18}O$ were measured simultaneously on a Picarro cavity ring-down spectroscopic analyzer, which interrogates water vapor directly through long path length absorption of nearinfrared energy emitted by a tunable diode laser system (www.picarro.com). Multiple replicates and internal standards are analyzed daily and calibrated against internationally recognized standards monthly. Precision of δD is ±1 per mil, and for $\delta^{18}O \pm 0.1$ per mil. The $\delta^{13}C$ sample is prepared by acidifying an aliquot of water under closed vacuum conditions to strip dissolved inorganic carbon from the water as carbon dioxide. The carbon dioxide is purified and then run separately on a dual inlet isotope ratio mass spectrometer that sequentially measures ${}^{13}C/{}^{12}C$ ratios of the sample and an internal standard carbon dioxide gas. Precision of the analysis is ±0.1 per mil. Note that for those samples with a suspended sediment load in the bottle, the sample was filtered in the lab prior to acidification to prevent isotopic bias from any carbonate sediments.

The ³H was measured by scintillation counting of the natural beta particle radioactive decay following electrolytic enrichment. This enrichment entails placing a cathode and anode in the water along with a strong electrolyte. The sample is kept sealed from the atmosphere while a current is induced through the water causing water molecules to dissociate into hydrogen and oxygen. The hydrogen gas that evolves has significantly less ³H than the remaining water, resulting in a stronger beta counting signal. Samples with electrical conductivities exceeding approximately 5,000 microsiemens (uS) required vacuum distillation prior to enrichment. The ¹⁴C was measured by first extracting carbon dioxide from a water sample using the same method as for ¹³C discussed above. The purified carbon dioxide was sealed in a glass ampule and shipped to Beta Analytic radiocarbon dating facility (www.radiocarbon.com). The ampule is opened under vacuum and reacted with hydrogen gas in the presence of an iron catalyst and converted to graphite. The graphite is packed into an aluminum target and placed on the source end of an accelerator mass spectrometer. The graphite is sputtered in the source, releasing ionized carbon atoms. The atoms are accelerated and separated by isotope by an electromagnet, and the ¹⁴C atoms are counted in a gas-filled detector. The ¹⁴C atom abundance is compared to standards and converted to a percent modern carbon (pmc). The pmc is percentage of the carbon relative to atmospheric carbon dioxide ¹⁴C abundance in ca. 1950. prior to any nuclear weapons testing influences. A ¹⁴C pmc of ≥100 is considered modern. A pmc <100 can be converted to an apparent age using the radioactive decay equation

$$Age = -8267 ln(^{14}C/^{14}C_{i})$$

where age is in years, ¹⁴C is the pmc of the sample, and ¹⁴C_i is the initial pmc at the time the sample was initially isolated from the modern carbon input and begins radioactive decay. In the case of an apparent age, ¹⁴C_i is 100 pmc.

For noble gas measurements, the copper tubes were fitted to an evacuated container under high vacuum. The copper cold seal formed during sampling is uncrimped and the water sample is released into the evacuated container where the water is subsequently degassed. A portion of the gas is separated and individual gas abundances are measured on a quadrupole mass spectrometer. The remaining gas is reacted with catalysts to remove nitrogen, oxygen, carbon dioxide, and hydrocarbon gases. The remainder is the noble gases. The isotopes of helium are measured specifically in this remaining gas on a magnetic sector mass spectrometer and detected using a Faraday cup detector. Results are reported as cubic centimeters of a specific noble gas per liter of water (ccSTP/g). The helium isotopes are also reported as a ratio of ${}^{3}\text{He}/{}^{4}\text{He}$ relative to the ${}^{3}\text{He}/{}^{4}\text{He}$ ratio in air. This ratio provides an indication of whether there is an excess of either ${}^{3}\text{He}$ or ${}^{4}\text{He}$ relative air. The dissolved noble gases are reported in addition as an amount of excess air in a sample (as excess neon, or $\% \Delta \text{Ne}$) and the calculated recharge temperature. The excess air results from entrainment of air during piston-type recharge, which dissolves air in excess of an equilibrium solubility.

6.0 Results

Samples were collected and analyzed for δD and $\delta^{18}O$ to delineate sources of recharge, and $\delta^{13}C$ to determine source of dissolved inorganic carbon. In addition, analysis was conducted on ¹⁴C and ³H to constrain the possible ages of groundwater recharge. Noble gas data measured at OLGEP monitoring wells serves the purpose of both a possible recharge source indicator and age constraint. Tabulated general water quality data is presented in **Table 3** and arranged into three groupings (Brine Pool, East Side, West Side) using a combination of pre-existing data and new data from this isotope study. Newly-acquired isotope data and general water quality data collected as part of this isotope study are shown on **Table 4**. Actual laboratory deliverables are included in **Appendix B** along with chain-of-custody documentation.

On **Table 4**, data are categorized into three groups comprising data from:

- 1) Springs and wells collected on the west side of Owens Lake,
- 2) OLGEP monitoring wells, and
- 3) Springs and wells on the east side of Owens Lake.

Sample sites were separated into these three categories based on the a) type of sample site, b) geographic area, and c) their relative differences in water quality as determined from field measurements. For example, on the west side, samples consisted of surface water, springs, and shallow wells. In contrast, the OLGEP monitoring wells sampled in the isotope study are nested, multi-level wells that are all artesian. East side samples included both shallow wells and springs, but measured field electrical conductivities were typically >1,000 uS compared to west side samples, which were below 1,000 uS. Furthermore, the isotopic data supports separation of the west side samples from the other two categories.

6.1 General Water Quality

Average concentrations of major cations and anions along with electrical conductivity and total dissolved solids are tabulated in **Table 3**. These data were compiled from existing water quality records derived from either LADWP and/or GBUAPCD and combined with data generated from the July 2012 sampling as part of the isotope study. The data are partitioned between east side and west side samples (as well as the brine pool itself) defined in general by sample location on either side of mapped faults along the west side of Owens Lake (**Figure 1**).

Sample Location Type	Average Ca (mg/L)	Average Mg (mg/L)	Average Na (mg/L)	Average K (mg/L)	Average Alkalinity (mg/L)	Average Cl (mg/L)	Average SO₄ (mg/L)	Average Electrical Conductivity (uS)	Average TDS (mg/L)	No. of Samples
			-		Brine Pool			-		
Brine Pool			116,800	4,840	89,600	120,600	25,200			5
					East Side					
Piezometers	6	26	15,480	488	12,088	13,715	3,467	48,975		4
Springs	22	52	1,099	67	1,425	840	231	4,998		12
OLGEP Wells		66	885		1,534	482		4,182	2,577	22
Other Wells	25	51	865	39	1,355	521	80	3,797		18
					West Side					
OLGEP Wells		2	739		929	534		3,024	1,775	6
Springs	19	5	145	6	261	68	37	769	254	6
Other Wells	14	2	24	3	76	13	13	200	105	5

Table 3Tabulated General Water Quality Data

Table 4 Tabulated General Mineral and Isotope Data Measured on Samples Collected in July 2012

Field Sample Name	Sampling Date	Field Temp. (°C)	Field Electrcal Conductivity (uS)	Field pH	Field Dissolved Oxygen (mg/L)	TDS (mg/L)	Ca (mg/L)	Na (mg/L)	Mg (mg/L)	K (mg/L)	Alkalinity (mg/L)	HCO ₃ (mg/L)	CO ₃ (mg/L)	SO₄ (mg/L)	CI (mg/L)	δD H₂O (per mil)	δ ¹⁸ O H ₂ O (per mil)	Tritium (TU)	Std. Dev.	δ ¹³ C DIC (per mil)	¹⁴ C DIC (pmc)	Std. Dev.	N ₂ total (ccSTP/g)	Ar total (ccSTP/g)	Ne total (ccSTP/g)	Kr total (ccSTP/g)	Xe total (ccSTP/g)	⁴ He (ccSTP/g)	R/Ra	Rech. Temp (°C)	∆Ne (%)	Notes
West Side	Jest Side																															
Carroll Crk	7/23/2012	15.6	133.0	8.4	9.1	69.0	12.8	4.9	1.5	0.6	42.0	42.0	0.0	10.7	7.2	-120.5	-15.94															
NW Seep	7/24/2012	20.9	882.0	6.4	0.5	409.0	28.2	116.0	5.5	5.4	178.0	178.0	0.0	48.1	49.2	-121.4	-15.80	< 1.00		-8.9	57.5	0.2										
Bartlett-CW	7/24/2012	22.2	517.0	8.4	3.7											-123.3	-16.16	< 1.00		-9.3	52.8	0.2										
Cottonwood Spr	7/24/2012	21.5	467.0	7.3	0.0	230.0	26.5	29.0	7.2	3.6	70.0	70.0	0.0	23.9	42.3	-120.4	-15.47	7.46	0.20	-15.7	88.9	0.3										
Cottonwood-Well	7/24/2012	23.5	6030.0	9.1	0.4											-103.5	-12.01															
Aqueduct Well	7/24/2012	14.5	97.0	5.9	6.7	56.0	7.2	4.2	1.8	1.4	30.0	30.0	0.0	9.9	7.1	-119.3	-15.40	3.82	0.17	-15.1	106.5	0.4										
Ash Crk	7/24/2012	19.6	260.0	7.9	8.3	122.0	23.7	8.0	3.7	1.9	60.0	60.0	0.0	29.0	7.4	-111.2	-14.56															
Rio Tinto-SW	7/24/2012	19.5	301.0	7.7	0.0	158.0	17.0	25.9	2.1	3.3	94.0	94.0	0.0	15.7	7.9	-114.5	-14.88	< 1.00		-9.8	56.1	0.2										
Cartago-COM	7/23/2012	18.0	253.0	8.3	6.3	138.0	24.4	12.1	2.1	2.0	80.0	80.0	0.0	14.6	7.4	-112.1	-14.71	1.76	0.15	-10.4	61.3	0.3										
S31	7/24/2012	20.0	1100.0	9.4	0.6	555.0	1.7	179.0	0.1	15.8	150.0	90.0	60.0	79.5	92.8	-110.2	-14.28	< 1.00		-3.7	25.5	0.2										
Duck-CN	7/24/2012	21.1	900.0	9.2	1.6											-111.7	-14.82															
OLGEP Monitoring	Wells																															
DWP-9 / T 896	7/23/2012	19.2	1460.0	9.2	0.0											-129.5	-16.55	< 1.00		21.7	2.9	0.1	1.05E-02	2.91E-04	1.27E-07	2.84E-08	1.28E-08	1.69E-05	1.52	-	-19.46	Large amount of excess He; poor gas model fit; sample stripped
DWP-9 / T 897	7/23/2012	19.3	1770.0	9.3	0.0											-131.8	-16.96															
DWP-9 / T 898	7/23/2012	19.9	1340.0	9.3	0.0											-129.4	-16.74	< 1.00		5.0	7.2	0.1	1.10E-02	3.83E-04	2.02E-07	1.12E-07	2.06E-08	4.38E-06	0.88	-	18.30	Large amount of excess He; poor gas model fit
DWP-3 / T 899	7/23/2012	19.4	1250.0	9.0	0.0											-123.3	-15.68	< 1.00		-1.5	3.8	0.1	1.16E-02	3.17E-04	1.84E-07	4.41E-08	1.17E-08	1.04E-05	0.91	-	5.23	Large amount of excess He; poor gas model fit
DWP-3 / T 901	7/23/2012	19.8	3570.0	9.3	0.0											-125.8	-16.11	< 1.00		0.7	1.8	0.1	8.60E-02	1.18E-03	3.12E-06	1.26E-07	1.33E-08	3.59E-06	1.17	-	1726.90	Large amount of excess He; poor gas model fit; excess air
DWP-6 / T 912	7/23/2012	20.0	5580.0	9.2	0.0											-132.7	-16.59	< 1.00		-0.2	1.7	0.1										
DWP-6 / T 911	7/23/2012	19.0	9820.0	8.8	0.0											-114.2	-13.25	< 1.00		-7.7	1.2	0.1	3.78E-02	3.15E-04	1.57E-07	7.00E-08	1.45E-08	7.84E-05	1.26	-	-9.37	Large amount of excess He; poor gas model fit; sample stripped
DWP-6 / T 913	7/23/2012	20.4	38900.0	9.2	0.0											-88.0	-8.28	< 1.00		0.9	0.5	0.1										
DWP-7 / T 908	7/23/2012	19.0	1680.0	9.2	0.0											-108.9	-14.34	< 1.00		-8.1	4.4	0.1										
DWP-7 / T 909	7/23/2012	19.1	747.0	9.3	0.0											-109.0	-14.38															
DWP-7 / T 910	7/23/2012	18.8	401.0	9.0	0.0											-110.1	-14.62	< 1.00		-6.0	36.7	0.2	1.21E-02	3.68E-04	2.29E-07	3.12E-08	1.04E-08	5.63E-07	0.81	16.9	36.07	Large amount of excess He; OK gas model fit
East Side																																
Dolomite	7/23/2012	25.1	1400.0	7.6	0.0											-127.6	-16.29	< 1.00		-7.6	5.6	0.1										
Agr-2	7/23/2012	20.3	1650.0	7.6	5.6											-125.0	-16.13	< 1.00		-5.8	3.6	0.1										
KCSD-PW	7/23/2012	22.3	1590.0	7.6	0.0											-124.5	-16.09	< 1.00		-5.5	4.4	0.1										
Horse Pasture	7/23/2012	22.1	2910.0	7.8	0.0											-130.9	-16.70	< 1.00		-5.1	1.9	0.1										
Tubman-CP	7/24/2012	26.4	8760.0	7.9	2.8											-128.8	-15.66															

With the exception of Na, cation data was generally lacking for several types of samples, which makes their comparison difficult. Anion data for alkalinity and Cl were more complete. In general, available data shows a systematic increase in Na, Cl and alkalinity from west side to east side samples. The highly evaporated brine pool shows the highest concentration, followed by east side piezometers and springs. East side wells are lower in these constituent concentrations. This general trend in east side samples is consistent with a process of relatively low salinity groundwater from depth (as recorded in wells) discharging to shallow springs and piezometers. During this discharge process, they are subjected to evaporative enrichment of their dissolved salts.

Comparison of the east side water chemistry to west side water chemistry shows much lower Na, Cl, and alkalinity concentrations, with concentrations of these analytes from wells being the lowest. This is consistent with the concept that west side groundwater is not being influenced by higher salt loads inherent in the subsurface below Owens Lake, and it implies that for the most part, groundwater beneath Owens Lake does not influence groundwater in west side wells and springs.

In **Table 4**, the OLGEP monitoring wells sampled in this study are treated as a third category because they penetrate well characterized hydrogeologic layers and by virtue of their exclusively artesian character. Because the general water quality data suggest distinct water quality types based on geography, the goal of the isotope data results is to determine if these distinct types are related to different recharge areas and different groundwater residence times.

6.2 Stable Isotopes of δD and $\delta^{18}O$

The δD - $\delta^{18}O$ values were measured on a total of 27 samples collected in the OLGEP study area. Values range from -16.96 to -8.28 per mil in δ^{18} O, and -132.7 to -88.0 per mil δ D. This range is consistent with values generated and reported in previous studies in the Owens Valley region (MWH, 2004). The δD and $\delta^{18}O$ data collected in this study are plotted in **Figure 2**. Also plotted for reference on this figure is the Global Meteoric Water Line (GMWL). This line is a linear fit to $\delta D - \delta^{18} O$ values of precipitation collected through different parts of the world. Most unevaporated precipitation should conform to this line. The OLGEP $\delta D - \delta^{18}O$ data are shown segmented into the three groupings in Figure 2. Note that the west side samples (red circles) partition into two distinct zones. The west side zone with higher isotopic values is all the sample sites located south of LADWP's Cottonwood Plant Well, whereas the lower isotopic zone includes samples collected north of Ash Creek. This division generally corresponds to a significant change in watershed source water. To the south, the highest elevations of the watershed are generally less than 10,000 feet above mean sea level (fmsl), whereas north of Ash Creek upper elevations draining into Carroll and Cottonwood creeks exceeds 10,000 fmsl. Recall as a general rule, higher elevation precipitation typically is associated with lower stable isotope values (see Appendix C).

One sample from the west side (Cottonwood Spring Well, adjacent to Cottonwood Spring) has a much higher isotopic value and lies further to the right of the GMWL. This position relative to the GMWL is consistent with water that has undergone significant evaporation that is believed to occur in the open standpipe of the abandoned well, which was not possible to purge prior to sampling. It is believed, therefore, that this isotopic value is not representative of underlying groundwater.

Isotopic values of OLGEP monitoring well samples shown in **Figure 2** (black circles) are predominantly lower than those of west side samples, with the exception of DWP-7. Values from DWP-7 are consistent with west side values south of LADWP Cottonwood Plant Well. This is not surprising because DWP-7 is geographically located in the same region as these west side samples.

In contrast, the three sampled monitoring wells from site DWP-9 wells had some of the lowest isotopic values. Isotopic values for the intermediate level of site DWP-6 wells (T912 - shallow portion of aquifer 5) were low similar to DWP-9. However, the deep and shallow at this site (aquifer 1 and deeper portion of aquifer 5) have significantly higher values that together form a linear evaporation slope projecting back to the intermediate level values. This indicates that not only is there evaporated water residing in the upper and lower sample levels of this well cluster, but also that the evaporated water evolved exclusively from the water in the intermediate level. This is why all three isotopic values line up along a straight line in **Figure 2**, consistent with evaporation. The isotopic values of the two levels sampled at site DWP-3 are somewhat higher than values at site DWP-9, and the intermediate level of site DWP-6. The isotopic values of the two levels sampled at site DWP-3 are somewhat higher than values at site DWP-9 and the intermediate level at site DWP-6. The deeper level of DWP-3 was highest. Note that both levels of site DWP-3 plot just below the west side cluster comprised of Carroll Creek, Northwest Seep, Bartlett Well, Cottonwood Spring NW, and LADWP Cottonwood Plant Well, whereas DWP-9 and DWP-6 are much lower. This does not imply that DWP-3 water is recharged from the west side, but rather it is likely recharged in part from a source with similar isotopic values, possibly derived from local east side recharge.



Note: The δD and $\delta^{18}O$ values for samples collected in July 2012 show distinct populations of surface and groundwater separating those sampled on the west side of Owens Lake area from those of the OLGEP monitoring well sites DWP-9, DWP-3, and DWP-6, as well as springs and groundwater sampled on the east side.

The isotopic values of the east side samples in **Figure 2** (blue squares) consistently overlap the lower isotopic values for the OLGEP wells. Their similarity suggests that they belong to the same population of groundwater in OLGEP monitoring wells with the exception of DWP-7.

The data plotted in **Figure 2** are also plotted against data from previous isotopic studies in **Figure 3**. Data from samples on the west side generally overlap with data presented in Lopes (1988) for samples that he collected on the west side. Likewise, the range in isotopic values of the OLGEP monitoring wells and east side samples are consistent with the range measured by the USGS throughout the Owens Valley. However, data from OLGEP monitoring wells and east side samples plot somewhat further to the right of the GMWL than the USGS data, suggesting that recharged groundwater under Owens Lake has undergone slightly more evaporation. Lastly, evaporated water in DWP-6 and Cottonwood Spring Well (adjacent to Cottonwood Spring NW), shown in **Figure 3**, overlap the region of evaporated water measured in piezometers presented in Font 1995 (data not shown).



Figure 3 $\delta D - \delta^{18} O$ Plot of OLGEP Data Compared to Previous Studies

Note: The range in isotopic values measured on samples collected in July 2012 generally overlap those presented in previous work by the USGS, Font (1995) (not shown), and Lopes (1988). OLGEP monitoring well sites DWP-3, and -6, as well as east side samples plot more to the right of the GMWL than USGS data, suggesting groundwater recharged under Owens Lake underwent slightly more evaporation.

Further consideration of the highly evaporated water residing within deeper aquifers sampled from DWP-6 is warranted. Care was taken to purge wells sufficiently before sampling to prevent any compromise of isotopic results. In addition, because the well is artesian, no open air exchange was possible to evaporate water within the well casing. However, an interesting comparison can be made between Owens River water and groundwater from DWP-6. **Figure 4** plots data from Coplen and Kendall (2000) that measured δD and $\delta^{18}O$ values of the Owens River at various times over nearly two years. Their samples were collected in the same spot each time below Tinemaha Reservoir near the town of Big Pine. The data form a general evaporation trend as would be expected from impounded surface water open to air exchange.

The slope of the Owens River evaporation trend is 5.37, suggesting that this is an average evaporative slope of surface water on a δD - $\delta^{18}O$ plot for this geographic region. Surprisingly, the linear fit of the three isotopic values of DWP-6 data is exactly the same (5.37), which is consistent with surface evaporation. This indicates that surface water that was significantly evaporated recharged aquifers intersecting DWP-6. The isotopic values of evaporated water in DWP-6 do not overlap with Owens River near Big Pine, and are therefore of a different source. The field electrical conductivity measurements are a further corroboration because they increase with increasing $\delta^{18}O$, consistent with evaporative enrichment of dissolved salts.



Figure 4 $\delta D - \delta^{18} O$ Plot Comparing Owens River to DWP-6

Note: Owens River water analyzed by USGS over various times of the year from 1984-1986 forms a surface water evaporation slope of 5.37. Surprisingly, a linear fit of the DWP-6 samples forms the exactly same slope, consistent with surface water evaporation, but from a water of a different source.

6.3 Isotopes of Carbon (¹⁴C and δ^{13} C) and Tritium

A total of 20 samples were measured for ¹⁴C, δ^{13} C, and ³H. The ¹⁴C values range from <1 to >100 percent modern carbon (pmc). The pmc value is a measure of the relative percent of the modern atmospheric carbon dioxide incorporated in the sample. Values greater than 100 pmc record abnormal ¹⁴C amounts introduced into groundwater recharge from surface testing of nuclear weapons. Values lower than 100 pmc either have undergone radiometric decay and/or have incorporated dissolved carbon from a ¹⁴C-absent source. Values at or below 1 pmc are essentially too low for practical groundwater age dating. **Figure 5** shows the distribution of ¹⁴C values are associated with the west side samples, ranging from >100 to 26 pmc. The OLGEP monitoring wells and east side samples on the other hand are all below 10 pmc, with the exception of the shallow level of DWP-7 (T910) at 37 pmc.

Tritium levels in all but 3 of the 20 samples were below detection limit of 1 tritium unit (TU). This is equivalent to one tritium atom per 10¹⁸ hydrogen atoms. The 3 samples where ³H was present were Cottonwood Spring NW, LADWP Cottonwood Plant well, and Cartago CSD. These 3 samples also had the highest ¹⁴C values. In general, the presence of ³H in a groundwater is usually an indicator of recharge, at least in part, of precipitation <50 years old, which is the time when surface testing of nuclear weapons occurred, and then left an annually decreasing residual ³H in the upper atmosphere.

The δ^{13} C values have a wide distribution, ranging from -15.7 to +21 per mil. The lowest value is consistent with dissolved inorganic carbon of shallow groundwater in isotopic equilibrium with soil carbon dioxide. This low value, as well as that of -15.1 per mil, is associated with samples Cottonwood Spring NW and LADWP Cottonwood Plant Well. The highest δ^{13} C values were measured for well DWP-9/T896 (+21 per mil) and DWP-9/T898 (+5 per mil). These high values are unusual for groundwater, which more often has an observed high value around 0 per mil, particularly for aquifers formed from carbonate rocks. However, a high value such as +21 per mil is not uncommon for dissolved inorganic carbon measured in oil and gas fields where methane is produced.

In DWP-9 samples, the δ^{13} C values are unusually high, indicating a carbon source distinct from sources contributing to the other wells.



Figure 5 ¹⁴C Comparison of OLGEP Samples

Note: The ¹⁴C values are highest in the west side samples and contrast those of the OLGEP monitoring wells (except for DWP-7/T910) and the east side samples. Only 3 samples had detectable ³H, which were Cottonwood Spring NW, LADWP Cottonwood Plant well, and Cartago-CSD.

The relationship between ¹⁴C and δ^{13} C is shown in **Figure 6**. It is interesting to note that a linear correlation occurs among the west side samples, as well as DWP-7/T910. The upper end-member of this correlation is Cottonwood Spring NW at 89 pmc and -15.7 per mil. The line extrapolates to a lower end-member between 0 and 1 per mil and 0 pmc. This lower end-member, defined by the low ¹⁴C and high δ^{13} C values observed in wells DWP-6/T913 and DWP-3/T901. The remaining samples plot to the left and below the line, with the exception of DWP-9, which plots to the right of the correlation. Samples that conform to this line could only a have undergone ¹⁴C reduction in proportion to δ^{13} C increases, indicating that no radiometric decay has taken place and only reaction with a ¹⁴C-absent, higher δ^{13} C source has occurred. Samples that plot below this line have likely undergone some radiometric decay, but the magnitude of decay cannot be determined simply from this plot because the extent of reaction with a ¹⁴C-absent source is not known.



¹⁴C- δ^{13} C Plot of Samples Collected

Note: West side samples and DWP-7/T910 form a linear correlation between δ^{13} C and 14 C with the upper end-member defined by Cottonwood Spring NW and the lower extrapolated end-member at δ^{13} C between 0 and 1 per mil and 14 C at 0 pmc. Some OLGEP monitoring wells conform to this lower end-member while others plot to the left and below the correlation. DWP-9 δ^{13} C values are consistent with a dissolved inorganic carbon in isotopic equilibrium with methane.

6.4 Noble Gas Results

The dissolved noble gases of helium (He), neon (Ne), argon (Ar), krypton (Kr), and xenon (Xe) were measured in groundwater on six individual samples collected at 4 separate OLGEP monitoring well sites. In addition, dissolved nitrogen was also measured. Nitrogen is a general measure of the amount of air dissolved in a groundwater. However, Ne is a better measure of dissolved air because nitrogen can periodically undergo reaction, such as in groundwater lacking dissolved oxygen. Consequently Ne is used in this data set as a quantitative measure of dissolved air. The other feature of Ne (and He) is that its solubility in water is relatively insensitive to temperature compared to Kr and Xe, which can change significantly with temperature (see **Appendix C**).

An approach then is to quantitatively gauge the amount of air dissolved in excess of equilibrium solubility in the water using Ne. Kr and Xe then is used to determine recharge temperature. The ⁴He concentration is used to estimate groundwater age based on its steady-state accumulation from crustal sources. The ³He is typically used for measuring the amount of ³H that has decayed since recharge.

Noble gas concentrations are reported as cubic centimeters of gas at standard temperature (20°C) and pressure (1 atmosphere) per gram of water, or ccSTP/g. The ccSTP is a volume of gas that would exist in air.

The Ne concentrations range from 1.27e-7 ccSTP/g in DWP-9/T896 to 3.12e-6 ccSTP/g in DWP-3/T901. This range is unusual for Ne because the solubility of Ne in water at STP is approximately 1.8e-7 ccSTP/g. The lowest concentration of 1.27e-7 ccSTP/g, as well as the concentration in DWP-6/T911 at 1.57e-7 ccSTP/g, suggests Ne is below equilibrium solubility, which is difficult to achieve in natural recharge conditions. The remainder of the samples exceeds equilibrium solubility. This is not unusual for groundwater, which often incorporates excess air during recharge. However, for sample DWP-3/T901, the excess amount is quite unreasonable and suggests entrainment of air by other means. This sample was noted for a high level of free gas in the water, and it was difficult to remove during the copper tube sampling. Consequently, given the abnormally high Ne concentration, it is likely that air bubbles were entrained in this sample, rendering the data unusable for further analysis. **Table 4** calculates the percent difference in Ne ($\%\Delta$ Ne) relative to equilibrium solubility to provide an easy method for assessing the amount of excess air. For the OLGEP monitoring well samples, only DWP-3/T899, DWP-7/T910, and DWP-9/T898 have reasonable excess air levels.

The He concentration measured as ⁴He in the samples ranged from 5.63e-7 ccSTP/g in DWP-7/T910 to 7.84e-5 ccSTP/g in DWP-6/T911. The equilibrium solubility of ⁴He at STP is approximately 4.9e-8 ccSTP/g, which indicates that OLGEP groundwater samples have excess ⁴He amounts ranging from about 10 to >1000 times equilibrium solubility. This excess amount of ⁴He far exceeds the excess amount of Ne measured in some of the samples, indicating that the ⁴He is from another source.

The ³He can be used as a ratio to ⁴He to indicate the source of the excess ⁴He. The ³He/⁴He ratio is shown as R/Ra in **Table 4**, which is the ³He/⁴He ratio of the sample normalized to the ³He/⁴He ratio of air. Note that equilibrium solubility concentration of 3He from air is approximately 7e-14 ccSTP/g. Helium derived simply from dissolution of air, R/Ra will be 1.0. The R/Ra will be <1.0, when He accumulates from crustal rocks only (alpha decay products). However, when the He is derived in part from active volcanic gas emission in the subsurface, R/Ra will be >1.0. The R/Ra of the OLGEP monitoring wells ranged from 0.81 to 1.52, suggesting that the excessive amounts of He dissolved in the groundwater are not solely from crustal accumulation since R/Ra would have been <<1.0. Consequently, it is highly probable that the excess dissolved He is from volcanic gas emission in the subsurface with R/Ra values >1.0. It may also be possible that active faults below Owens Lake are sources of deep He emissions.

Unfortunately, the Ne in some samples was below equilibrium solubility, and He was dissolved in excessive amounts. Normally Ne will always be at or above equilibrium solubility, and ⁴He concentration will be proportional to excess air amounts plus any small amount accumulated from crustal emission. In the OLGEP monitoring wells, ⁴He far exceeds normal crustal accumulated amount. Furthermore, because Ne is below solubility in some samples, excess air calculations are not possible. This also makes the remaining samples suspect that some Ne loss may have occurred as well. Most of the OLGEP monitoring wells sampled had some amount of free gas in the water during sampling. Anecdotal evidence suggests this free gas is methane. Excess amounts of methane in the groundwater could be preferentially removing Ne and Ar over Kr and Xe during its ascent up the well bore. Ne and Ar are less soluble than Kr and

Xe and would be remove more readily. Consequently, it is probable that the free gas in these wells have disrupted the relative abundance of noble gases measured.

Only in the case of DWP-7/T910 is the excess He and excess Ne reasonable enough that excess air can be determined and recharge temperature calculated. The recharge temperature reported is 16.9°C (**Table 4**). The measured water temperature during sampling was 18.8°C. The implications of this calculated recharge temperature are discussed in section 7.1.

7.0 Update of Interpretive Model

In the SAP (MWH, 2012), general interpretive conclusions were drawn from previous isotope studies in the Owens Valley. Early work by Lopes (1988) suggested that $\delta D - \delta^{18}O$ measurements of springs and wells around the perimeter of Owens Lake showed differences that likely were related to different recharge sources. Font (1995) measured $\delta D - \delta^{18}O$ in shallow piezometers installed around the perimeter of Owens Lake and showed that much of the shallow groundwater was evaporated. The slope of the evaporation line on a $\delta D - \delta^{18}O$ plot was approximately 4.3, consistent with shallow subsurface evaporation. The data all seemed to be extrapolating back to a fairly uniform source consistent with deeper groundwater in the eastern portion of the lake measured by the USGS (Densmore et al., 2009).

Using ¹⁴C and δ^{13} C data, previous studies further determined that there appeared to be a fairly uniform reaction between young recharge (with higher ¹⁴C and low δ^{13} C values) and a ¹⁴C absent source (with higher δ^{13} C values), and that this reaction would largely control apparent groundwater age calculated from ¹⁴C measurements.

The SAP also defined three main goals of the sample collection program:

- 1) Determine how uniform groundwater isotopic signatures were in groundwater beneath Owens Lake,
- 2) Define the isotopic character of springs and shallow wells on the west side of Owens Lake and how they may be related to deeper groundwater beneath the lake, and
- 3) Determine the isotopic character of shallow groundwater on the east side of the lake.

This TM articulates findings relative to the goals of the study and further advances the interpretation of recharge sources and ages and how they inter-relate.

7.1 Sources of Recharge in the OLGEP Study Area

The variation in the δD - $\delta^{18}O$ values of surface and groundwater in the Owens Lake area serves as a foundation for interpreting sources of recharge. When the $\delta^{18}O$ values are viewed in map form and contoured, the variation reveals a systematic pattern that illustrates different sources contributing to recharge (**Figure 7**). For instance, the lowest $\delta^{18}O$ values (-16.5 per mil) occur as a tongue leading from north to south. These lower $\delta^{18}O$ becomes higher on the west, south, and east sides of Owens Lake. Increase in $\delta^{18}O$ is controlled by local recharge around the lake. **Figure 7** also includes data from the previous USGS study (Densmore et al., 2009).



This map has been designed to print size 8.5" by 11".

Note that the lowest δ^{18} O values in the map are controlled exclusively by OLGEP monitoring well water, with the exception of site DWP-7. All three depths sampled and measured at DWP-9 (located at the head of this tongue) show a relatively uniform δ^{18} O value. In contrast, DWP-3 on the eastern side of the tongue, has a lower value at its shallow depth (T901: 150-170 foot depth screened in Aquifer 1), but somewhat higher at its deepest (T899: 920-960 foot depth screened in Aquifer 5). The three sampled depths at DWP-6, located on the south end of the tongue, have large δ^{18} O variations due to recharge of highly evaporated water from a single source. Only the mid-level has been used in **Figure 7** to avoid weighting of the δ^{18} O tongue and the terminus of the modern Owens River. The low δ^{18} O values in this groundwater are unrelated to any surface water and shallow groundwater in the immediate surroundings of Owens Lake based on data presented above. Consequently, modern recharge sources for this low δ^{18} O water defines the down-valley flow component characterized in the Updated OLGEP Conceptual Model Report (MWH, 2011).

One alternative interpretation of the low δ^{18} O values seen in OLGEP monitoring wells is the possibility that they are an isolated pocket of pluvial period recharge. In other words, they are ancestral recharge, or connate waters remaining from a different historical climatic regime. This interpretation is commonly forwarded where groundwater with isotopic values that are appreciably lower than shallow groundwater locally recharged . This interpretation is commonly wells and when all other potential sources can be ruled out. However, the interpretation is commonly employed without thorough analysis, and the research literature is fraught with such erroneous results (see Davisson et al., 1999 for further discussion). In the OLGEP study area, numerous upgradient recharge sources with the low δ^{18} O values are possible as source water for low δ^{18} O groundwater in OLGEP monitoring wells.

Recall that the δD - $\delta^{18}O$ values of surface water and shallow groundwater along the western side of Owens Lake suggest that mean elevation of upgradient runoff control their variation, illustrated by distinct isotopic populations north and south of Ash Creek (**Figure 2**). The lower $\delta^{18}O$ values of De La Cour Ranch/Carroll Creek, Northwest Seep, and groundwater in the Alabama Hills (from USGS study data) are not sufficiently low enough to be a reasonable recharge source for the low $\delta^{18}O$ groundwater beneath Owens Lake, as observed at site DWP-9, in particular. Further north of the OLGEP study area, even lower isotopic values occur in runoff and shallow recharge, as shown by previous isotopic measurements in the Inyo/LA Cooperative Geochemical Study (MWH, 2004). Results of that study found that surface water and shallow groundwater on the west side of Owens Valley consistently showed $\delta^{18}O$ values ranging less than -17 per mil. It is likely that mean $\delta^{18}O$ of runoff from the Sierran watersheds decreases from south to north along Owens Valley. This decrease provides some geographic control on defining recharge sources in areas where mixtures of local and more regional groundwater occur.

Another factor in considering an upgradient recharge source is the fact that the lowest δ^{18} O groundwater beneath Owens Lake also has significantly lower ¹⁴C and higher δ^{13} C values. Regardless of whether this is due mostly from radiometric decay or from contribution of ¹⁴C-absent sources, data indicates that this groundwater has been isolated sufficiently long enough to have undergone gross geochemical changes compared to groundwater on the west side of Owens Lake.

The distribution of low δ^{18} O water on the east side of this tongue is somewhat complicated by the fact that at site DWP-3 shallow (T901) (-16.1 per mil) and deep (T899) levels (-15.7 per mil) are different (these values were arithmetically averaged for the contour map). The deeper, higher δ^{18} O water may be due to infiltration of more local recharge in this area. However, no further conclusions can be drawn about local recharge here except that the higher δ^{18} O water is similar to that found further south and northwest side of Owens Lake.

One of the most important features is the general correspondence between the distribution of the low δ^{18} O groundwater tongue and the position of mapped faults. This correspondence is most remarkable where faulting in the southwestern portion of Owens Lake occurs between OLGEP monitoring well sites DWP-6 and DWP-7. Recharge sources of these two sites are distinctly different. DWP-7 is clearly recharged from local sources (similar to Ash Creek) that also recharge wells Rio Tinto, Cartago-CSD, S3(1), and Duck 3. DWP-6, on the other hand, has an unevaporated signature similar to groundwater in wells at site DWP-9. It is reasonable to assume that the northward extension of this fault group in the western portion of Owens Lake also contributes to the segmenting of western shallow recharge from deeper groundwater beneath Owens Lake.

On the east side of Owens Lake, correspondence between faulting and δ^{18} O values of groundwater is less certain. Much of this uncertainty may be due to lack of enough shallow collection sites and data on the east side compared to the type of samples on the west. Nevertheless, the fact that groundwater at site DWP-3 did show isotopic layering between one depth and another might suggest a lack of fault control and more control by hydraulic conductivity. Until further data can be generated, east side fault control of groundwater flow is uncertain at this time.

It is useful to discuss the recharge temperature that was successfully calculated for OLGEP well DWP-7/T910 from the noble gas results. The calculated recharge temperature was 16.9°C. In order to assess this calculated temperature and its implication for recharge source, a temperature/elevation curve needs to be constructed in this general region of eastern California. This can be accomplished by plotting mean annual air temperatures measured at various meteorological stations in California. This was done for 14 different stations, within or close to Owens Valley, ranging from <2,000 to >12,000 fmsl (**Appendix D** and **Figure 8**). A linear regression was fit to the data.

Mean air temperatures of 14 meteorological stations at various elevation are correlated to yield a linear regression and slope that indicates approximately 2°C change in temperature occurs for every 1,000 feet of elevation change. Data were obtained through the Western Regional Climate Center (http://www.wrcc.dri.edu/summary/Climsmcca.html) and are tabulated in **Appendix D**.



Elevation-Mean Air Temperature Plot for the OLGEP Region

This linear regression is used to compute an elevation from the noble gas recharge temperature. The calculated elevation is approximately 2,700 fmsl using the regression in **Figure 8**. Unfortunately, this elevation is below the ground surface of Owens Lake and suggests an additional consideration is needed in order to derive a recharge elevation. One issue that has been discussed at length in research literature is differences between soil temperature near the water table and mean air temperature (e.g., Stute and Sonntag, 1992). Diurnal, as well as seasonal changes in air temperatures, do not influence water table temperature that occurs at depth where the noble gas recharge temperature is set. However, this temperature at depth is typically higher than the mean air temperature, between 2 and 4°C. If we assume that it is 3°C for the recharge area of DWP-7/T910, then the mean air temperature above that recharge zone would be 13.9°C, instead of 16.9°C. The 13.9°C temperature is similar to 14.5°C temperature measured in LADWP Cottonwood Plant Well, which had the youngest groundwater age. The 13.9°C temperature places the calculated recharge elevation for DWP-7/T910 at approximately 4,150 fmsl using the regression in **Figure 8**.

Recharge along the west side of the OLGEP area is dominated by stream recharge. Infiltration is maximized in the springtime during intense snow melting. The infiltration occurs as streams emerge from exposed bedrock surfaces at elevation and begin to cross the porous alluvial fan deposits at lower elevations. The 4,150-foot elevation just west of DWP-7/T910 correlates well with the upper end of the alluvial fan systems and stream flow emerging from the higher bedrock elevations. Consequently, the calculated recharge temperature and recharge elevation

correlate with a plausible recharge are for DWP-7/T910. Furthermore, the δD - $\delta^{18}O$ values for DWP-7/T910 correlate to the west side recharge as shown in **Figure 2**.

7.2 Groundwater Age Interpretation

The ³H, ¹⁴C, and δ^{13} C results are most applicable to the west side groundwater. As shown is **Figure 5**, this groundwater has the highest ¹⁴C abundance, with three samples having measurable ³H. All of the west side samples, except for the two lower levels of OLGEP wells at site DWP-7, line up on a straight line in **Figure 6** that extrapolates between modern recharge (¹⁴C = 89-106 pmc) and a ¹⁴C-absent source with δ^{13} C between 0 and 1 per mil. This strongly suggests that this groundwater uniformly undergoes a fairly rapid reaction with a ¹⁴C-absent source. The reaction is rapid because there is no indication that this groundwater has experienced any radiometric decay of its ¹⁴C, which would plot data below the straight line in **Figure 6**. As further evidence that a single source is contributing to the decrease in ¹⁴C and increase in δ^{13} C, the alkalinity as milli-equivalents (meq) is plotted against the ¹⁴C and δ^{13} C separately in **Figures 9a and 9b**. Milliequivalent is calculated by dividing each constituent contributing to the alkalinity (HCO₃ and CO₃) by its molecular weight and multiplying by its charge state (times 2 in the case of CO₃). This provides a better measure of the dissolved carbon reactivity progress.

First of all, both plots show a linear array of data that have been extrapolated to end member values. In Figure 9a, the ¹⁴C extrapolates to an alkalinity between 4 and 5 meg at 0 pmc. At the other end at 0 meg, the ¹⁴C is between 100 and 120 pmc, consistent with atmospheric carbon dioxide values in the past 20 years. In Figure 9b, the extrapolation to an alkalinity between 4 and 5 meg results in a δ^{13} C values of approximately 0 per mil, a value consistent with marine carbonates. The other end of the extrapolation intersects a δ^{13} C value of -18 per mil at 0 meg. The -18 per mil value is consistent with a soil zone bicarbonate species in isotopic equilibrium with carbon dioxide of approximately -27 per mil at presumed soil temperatures in this area. The -27 per mil value is consistent with a respired carbon dioxide by plants utilizing a Calvin cycle metabolic pathway. The Calvin cycle is a type of biochemical reaction series that occurs in many common types of plants during photosynthesis. During this cycle, the plant will preferentially take up carbon dioxide with a low ¹³C content and respire excess carbon dioxide with the same low ¹³C abundance. The δ^{13} C value of this respired carbon dioxide ranges between -25 and -30 per mil. Dissolved bicarbonate in open exchange with a carbon dioxide of this isotopic range will subsequently take up a proportionally higher amount of ¹³C over ¹²C, and the difference in δ^{13} C between the bicarbonate and the carbon dioxide is generally between 8 and 10 per mil. This is why the -18 per mil extrapolated end-member in Figure 9b is consistent with bicarbonate in isotopic equilibrium with soil zone carbon dioxide. In summary then, the commonality between Figures 9a and 9b supports a uniform reaction process for west side samples.



Figure 9 Comparison of Alkalinity with ¹⁴C and δ^{13} C Values in West Side Wells

Note: Alkalinity in milli-equivalents for west side samples plotted against (a) ¹⁴C and (b) δ^{13} C show linear arrays that when fitted with a best fit line extrapolate to end-member values consistent with those end-members defined for west side samples in Figure 5. This supports the conclusion that a uniform ¹⁴C-absent source contributes to decreasing ¹⁴C and increasing δ^{13} C of west side groundwater.

The fact that there appears to be no radiometric decay of the ¹⁴C in the west side samples likely limits the age of these groundwater to <300 years. The 300-year limit is determined by the measurement precision (approximately ±1 pmc) compared to ¹⁴C decay is this timeframe. In the case of Cottonwood Spring NW, LADWP Cottonwood Plant Well, and Cartago-CSD samples, there is measureable ³H, which would limit these groundwaters to ages of <50 years old. It is further likely that these youngest groundwaters are less than 10 years old, since their ³H concentrations are too low to represent 1960s recharge.

The groundwater age determination for the OLGEP wells and east side samples is not as straightforward as the west side samples. This is due largely to the fact that:

- 1) No ³H is detected in the OLGEP wells or east side samples,
- 2) The ¹⁴C abundance is low, and
- 3) The δ^{13} C values tend to be enriched.

There also appears to be no systematic change in the ¹⁴C and δ^{13} C values as observed in the west side samples. Consequently, determining how much radiometric decay has occurred versus reaction with a ¹⁴C-absent source is very difficult, particularly since recharge sources and flow paths are not clearly delineated, and aquifer geochemistry is still poorly understood. Nevertheless, it is possible to calculate minimum and maximum ages for those samples that have detectable ¹⁴C and δ^{13} C values less than 0 per mil. This will exclude samples from OLGEP wells at site DWP-9 and the shallow levels at sites DWP-3 and DWP-6. The remaining samples can be treated by two different approaches. The first assumes that all the reaction with a ¹⁴C-

absent source occurred prior to radiometric decay. This duplicates what has been interpreted to be controlling the west side groundwater ¹⁴C values and uses the correlation line in **Figure 6**. It also will calculate the youngest possible ages for the OLGEP wells and east side samples. The second approach assumes all the radiometric decay occurred prior to δ^{13} C enrichment. This simply ignores the effect of subsurface reaction with a ¹⁴C-absent source. This is also known as an apparent age. Age calculations are shown below in **Table 5**.

Sample	δ^{13} C per mil	¹⁴ C pmc	Apparent Age (years)	Minimum Age (years)		
OLGEP Monito						
DWP-9 / T 896	21.7	2.9	29,269			
DWP-9 / T 898	5.0	7.2	21,751			
DWP-3 / T 899	-1.5	3.8	27,034	20,621		
DWP-3 / T 901	0.7	1.8	33,212			
DWP-6 / T 912	-0.2	1.7	33,684	28,017		
DWP-6 / T 911	-7.7	1.2	36,564	6,759		
DWP-6 / T 913	0.9	0.5	43,801			
DWP-7 / T 908	-8.1	4.4	25,823	6,958		
DWP-7 / T 910	-6.0	36.7	8,287	0		
East Side						
Dolomite	-7.6	5.6	23,829	7,739		
Agr-2	-5.8	3.6	27,481	9,538		
KCSD-PW	-5.5	4.4	25,823	10,195		
Horse Pasture	-5.1	1.9	32,765	10,096		

Table 5 Calculated Groundwater Ages in OLGEP Monitoring Wells and East Side Samples Using ^{14}C and $\delta^{13}\text{C}$

The calculated age range shows differences between minimum and maximum ages that are small, such as in DWP-3/T899, and large, such as seen DWP-6/T908. This approach to bounding groundwater ages makes relatively simple assumptions about the history of inorganic carbon chemistry in groundwater leading to observed isotopic values. Its history may be far more complex and requires more extensive correction, but at this time there is no supporting data to take this approach. For example, the abnormally high δ^{13} C values for DWP-9 imply that the inorganic carbon is greatly influenced by active methane generation in the subsurface. This production is likely biological and follows an acetate fermentation pathway. Acetate fermentation is the commonly observed process of complex sugar breakdown to carbon dioxide and methane by micro-organisms in oxygen-limited environments. Examples would include processes that generate carbon dioxide and methane within sanitary landfills, or the well-known fermentation process that leads to alcoholic beverages. This co-production of carbon dioxide and methane will result in an isotopic equilibrium being established between these two species. The partitioning or difference between these two species will be approximately 70 per mil δ^{13} C, with carbon dioxide being high and the methane low. Biogenic methane is typically less than -50 per mil, which will make the carbon dioxide as high as approximately +20 per mil, similar to that observed in DWP-9/T896. Consequently, the ¹⁴C abundance of the dissolved inorganic carbon
in this groundwater will be controlled mostly by the decaying organic matter and not the recharge source or other inorganic carbon sources.

The ⁴He results from the noble gas data unfortunately do not provide any additional insight into potential groundwater ages in the OLGEP wells. This is due to the excessive amounts of ⁴He observed above what is naturally diffused upward from the earth's crust. State-state accumulation of crustal ⁴He in groundwater averages around 5.0e-7 ccSTP/g for every 10,000 years. Comparison of this rate to the ⁴He concentrations in **Table 4** would indicate groundwater ages ranging from approximately 10,000 to 1.6 million years old. These ages are entirely inconsistent with both the hydrologic and isotopic evidence collected for OLGEP groundwater. We know little about the actual source area or the rate of high He discharge. Although it is likely related to both recent volcanic activity to the south in the Coso area, as well, as active faulting beneath Owens Lake, it is not known if it is uniformly diffuse through the basin, or discharges in concentrated areas.

8.0 Implications of Isotopic Sample Results on Estimated Aquifer Parameters

The results of isotopic analyses can be used to provide yet another method to estimate hydraulic conductivity based on estimated travel time (groundwater age) and generalized hydrogeologic assumptions. These estimations can then be compared to the current conceptual hydrogeologic model (MWH, 2011).

Using the west side ³H and ¹⁴C results, it is possible to produce a general estimate of aquifer hydraulic conductivity using a simple average linear groundwater velocity equation:

$$v = \frac{K}{n} \frac{dh}{dl}$$

where *V* is the average linear velocity, *K* is the hydraulic conductivity, *n* is the effective porosity, and dh/dl is the hydraulic gradient.

The linear velocity can be calculated for sample sites on the west side by taking the estimated age and dividing it by the distance to an assumed recharge area. If it is assumed that most of the west side recharge is primarily derived from the apex of the alluvial fans, then the furthest point of recharge would be the apex of the closest alluvial fan. This distance would provide the greatest length of groundwater travel. The hydraulic gradient then is estimated using the linear velocity equation coupled with the following generalized assumptions:

- Assign groundwater ages to be the youngest plausible (5 years for water with ³H, and 60 years for others),
- Assume that the hydraulic gradient is half that of the surface topography, and
- Assume an effective porosity of 0.2.

Results of this exercise are tabulated in **Table 6**. The estimated range in hydraulic conductivity is comparable to previous conceptual models, as reported in MWH (2011).

	Determinat	tions from ³ H	and ¹⁴ C	
Sample Location	Distance (miles)	Age (years)	dh/dl (-)	Estimated K (ft/day)
Cartago-CSD	1	5	0.025	23.15
Northwest Seep	3	60	0.025	5.79
Bartlett Well	3	60	0.025	5.79
Cottonwood Spring NW	2	5	0.025	46.29
LADWP Cottonwood Plant Well	0.5	5	0.025	11.57
Rio Tinto	0.75	60	0.025	1.45

60

60

0.025

0.025

5.79

3.86

Table 6Estimate of Aquifer Hydraulic Conductivities on the West Side of Owens Lake Using AgeDeterminations from ³H and ¹⁴C

9.0 Conclusions and Recommendations

S3(1)

DWP-7/T910

3

2

Water quality data and isotope results are both useful in elucidating sources of recharge and for groundwater age-dating in the OLGEP study area. Major cations and ions provide a simple measure of mixing between dilute and saline water sources. Stable isotopes of deuterium (D) and oxygen-18 (¹⁸O) provide a simple approach for delineating specific recharge areas, and tritium (³H) establishes whether a groundwater was recharged within the past 50 years. Radiocarbon (¹⁴C) proves less useful for age-dating because of a ubiquitous occurrence of ¹⁴C-absent minerals in the OLGEP study area that readily interact with dissolved inorganic carbon in groundwater. Active methane generation at depth below Owens Lake also interferes with ¹⁴C age-dating. Likewise, age-dating and recharge temperature calculation from measured dissolved noble gases (He, Ne, Ar, Kr, Xe) proved elusive for OLGEP wells because of an unusually high gas emission rate beneath the lake, in particular helium, that apparently disrupts the relative abundance of noble gases. However, in one well location (DWP-7/T910), a recharge temperature was successfully calculated and correlated well with the most plausible recharge area.

Specific conclusions from the data generated in this study are listed below:

- The total dissolved solids concentrations are low in groundwater and springs on the west side of Owens Lake, which contrasts with the more saline groundwater beneath Owens Lake at depth, as well as springs and shallow wells on the east side. The dilute concentration observed in west side samples implies that they are not under the immediate influence of the more saline water to the east.
- 2) The δD-δ¹⁸O results demonstrate that the dilute west side waters are recharged by stream channel infiltration originating from the higher elevations of the eastern Sierra Nevada. This is further supported by the similarity in isotope abundances between measured creek samples and local groundwater and springs.

- 4) The δD-δ¹⁸O of sampled OLGEP monitoring wells (except site DWP-7) demonstrates that they are not recharged locally, but rather their lower isotopic values are consistent with isotopic values of recharge sources further north. In other words, water in OLGEP monitoring wells is consistent with a down-valley flow source.
- 5) Based on the δD-δ¹⁸O results, it can be concluded that groundwater at OLGEP well site DWP-6 located in the southeastern portion of Owens Lake was recharged by significantly evaporated surface water. This suggests that historically, ponded water may have recharged groundwater aquifers locally in the southern portion of the Owens Lake area.
- 6) The δD - $\delta^{18}O$ values in groundwater at OLGEP well site DWP-7 indicates recharge from the west side.
- 7) Based on the geographic distribution of groundwater δ¹⁸O values in the OLGEP study area, mapped faults beneath the western side of Owens Lake are likely limiting groundwater movement between west side recharge and down-valley flow sources.
- 8) The δD-δ¹⁸O values for springs and shallow wells located on the east side of Owens Lake are generally the same as those for OLGEP monitoring wells recharged by downvalley flow, which indicate a similar isotopic and geochemical character. However, there are some differences observed in the isotopic values, such as in DWP-3/T899, that suggest locally-recharged groundwater does contribute to some recharge on the east side.
- 9) Groundwater ages measured for springs and shallow wells on the west side of Owens Lake were found to be less than 300 years old.
- 10) Groundwater from three sites (Cottonwood Spring NW, LADWP Cottonwood Plant Well, and Cartago-CSD) was found to be less than 50 years old.
- 11) Hydraulic conductivities calculated using isotopic results are comparable to previous conceptual models, as reported in MWH (2011), and appear reasonable.
- 12) Calculated ¹⁴C groundwater ages ranged from approximately 6,800 years to greater than 40,000 years in OLGEP monitoring wells that are recharged by down-valley flow, as well as sample sites on the east side of Owens Lake.
- 13) The ¹⁴C measured in OLGEP well site DWP-9 was heavily influenced by active methane production at depth, and age-dating is highly uncertain.
- 14) Uncertainties remain for groundwater ages in OLGEP wells and east side sample sites. The uncertainty results from:
 - a poor control on reaction mechanisms that cause ¹⁴C-absent minerals to contribute carbon to the dissolved inorganic carbon in groundwater.
 - an usually high helium gas emission occurs below Owens Lake that obscures the steady-state accumulation of ⁴He commonly used in semi-quantitative age-dating of groundwater.
- 15) The unusually high helium gas emission is also accompanied by a high discharge of other gas presumed to be methane and was observed in several of the OLGEP wells.
- 16) This high gas emission rate disrupts the relative abundance of the dissolved noble gases in the aquifer and the wells by preferentially stripping out neon and argon over krypton and xenon. This disruption has made the measured noble gas abundance inadequate for calculating ages and recharge temperatures.

17) Only in the case of OLGEP well DWP-7/T910 was it possible to calculate recharge temperature. The recharge temperature of groundwater in DWP-7/T910 was calculated at 16.9°C, which correlates with a recharge elevation of approximately 4,100 feet above sea level. This elevation is consistent with the upper end of alluvial fans west of DWP-7/T910, where stream channel infiltration takes place.

This limited isotopic study was successful in meeting the goals outlined in the SAP and improved the hydrogeologic conceptual model in several ways. Future isotopic and general water quality sampling and analysis may prove useful for shedding light in key areas:

- Previous sampling by GBUAPCD, LADWP, and others utilized varying sample suites for cations and anions. A uniform sampling and analysis plan for general minerals (cations and anions) across the study area could provide a useful baseline of water quality. General minerals could also be used as a simple monitoring tool for potential migration of saline water into less saline groundwater following a change in groundwater use patterns.
- 2) Better resolution for recharge sources of down-valley could be provided by additional measurements of δD and $\delta^{18}O$ conducted on a number of surface and groundwater sites in suspected recharge areas and along flow-paths leading to the OLGEP study area.
- 3) Better age-dating of down-valley flow may be provided by sampling for ¹⁴C and noble gases at OLGEP monitoring well sites DWP-1 and DWP-11 (or similarly accessible wells with well-known depths) and/or down-valley flow sites east of the Owens River where there is well construction data and control on the depth of sampling. However, the sampled wells should be absent of free gas of the type that was observed in site DWP-9.
- 4) Faulting in the eastern side of Owens Lake that influences groundwater migration could be further defined by additional samples collected for δD and δ¹⁸O, particularly for any wells that might be upgradient of the lake level or adjacent wells with distinctly different depths. This will help establish the influence of any locally recharged groundwater flowing towards Owens Lake, and whether this locally-recharged groundwater comingles with down-valley flow.
- 5) Because $\delta D \delta^{18}O$ measurements are diverse enough among the different recharge sources, additional measurements could be used as a simple and inexpensive groundwater recharge mapping tool.

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APPENDIX A

Site Summary Sheets

OLGEP Task:	401.1.9.2 - Perform	i Isotope San	npling and <i>i</i>	Analysis of Gr	oundwater and Su	ırface Discharge Ar	eas		
Site Name:	De La Cour Ranch	(Carroll Cree	k)		Field Label:	Carroll Crk			
Description of Site:	Surface Water from	n Carroll Cree	łk		Date Sampled:	7/23/2012			
	Northing:	4,040,470.7	6		Well Depth (ft):	t): N/A			
Coordinates:	Easting:	401,272.19			Screened	s)- N/A			
	Elevation (fmsl):	5,497			Intervals (fbgs):				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)	
	15.60	8.35	-65	121	0.13	0.0	9.05	91.9	
	Cations	5/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas	
Samples Submitted to Laboratory:	Anions	6	δD a	nd $\delta^{18}O$	δ^{13} C				
	X			x					





Location of Sample: Ponded Water on Carroll Creek Upstream of Culvert West of Horseshoe Meadows Road

Measurement of Field Parameters at Culvert on Carroll Creek





Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perfori	m Isotope Sar	npling and a	Analysis of Gr	oundwater and Su	urface Discharge Ar	eas	
Site Name:	Northwest Seep				Field Label:	NW Seep		
Description of Site:	Spring Water from	NW Seep Ar	ea		Date Sampled:	7/24/2012		
	Northing:	4,040,533.8	6		Well Depth (ft):	N/A		
Coordinates:	Easting:	407,810.09			Screened	N/A		
	Elevation (fmsl):	3,575			Intervals (fbgs):		Dissolved Oxygen (mg/l) 0.46 Tritium	
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)
	20.87	6.36	31	-9	0.88	280.0	0.46	4.7
	Catior	ns/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas
Samples Submitted	Anior	IS	δD a	nd $\delta^{18}O$	δ ¹³ C			
	x			x	x	x	x	





View to Southeast Showing Northwest Seep



Sample Preparations at Northwest Seep





Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform	n Isotope Sar	npling and <i>i</i>	Analysis of Gr	roundwater and Su	urface Discharge A	reas	
Site Name:	Bartlett Well				Field Label:	Bartlett-CW		
Description of Site:	Groundwater from	Bartlett Flowi	ng Well		Date Sampled:	7/24/2012		
	Northing:	4,037,918.2	7					
Coordinates:	Easting:	408,049.05			Screened	161 - 284	(based on 2008 v associated interp Dissolved Oxygen (mg/l) 3.66 Tritium	rideo log and retation)
	Elevation (fmsl):	3,582			Intervals (fbgs):	500 - 575		
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)
	22.15	8.35	-74	-78	0.52	0.0	3.66	41.9
	Cation	s/		Stable Isot	topes	Radiocarbon	Tritium	Noble Gas
Samples Submitted	Anion	S	δD a	nd $\delta^{18}O$	δ^{13} C			
				x	x	x	x	





View to the South of the Barlett Well Sample Location



Close-Up Photo of Barlett Well Showing Flowing Water





Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform	i Isotope San	npling and <i>i</i>	Analysis of Gr	oundwater and Su	ırface Discharge Ar	eas		
Site Name:	Cottonwood Spring	NW			Field Label:	Cottonwood Spring			
Description of Site:	Spring Water from	Cottonwood	Spring NW		Date Sampled:	7/24/2012			
	Northing:	4,032,155.6	3		Well Depth (ft):				
Coordinates:	Easting:	408,943.32			Screened N/A				
	Elevation (fmsl):	3,563			Intervals (fbgs):	npied: 7/24/2012 th (ft): N/A d (fbgs): N/A rical ctivity cm) Turbidity (NTU) Dissolved Oxygen (mg/l) Dissolved (% 17 371.0 0.00 0			
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)	
	21.52	7.33	-19	-72	0.47	371.0	0.00	0	
	Cations	5/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas	
Samples Submitted	Anions	6	δD a	nd δ ¹⁸ O	δ^{13} C				
	x			x	x	x	x		









Cottonwood Spring Dug Hole





Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform	Isotope San	npling and <i>i</i>	Analysis of Gr	oundwater and Su	ırface Discharge Ar	eas		
Site Name:	Groundwater from	Cottonwood S	Spring Well		Field Label:	Cottonwood-Well			
Description of Site:	Well located near (Cottonwood S	pring		Date Sampled:	7/24/2012			
	Northing:	4,032,155.63	3		Well Depth (ft):	(ft): Unknown			
Coordinates:	Easting:	408,943.32			Screened				
	Elevation (fmsl):	3,566			Intervals (fbgs):	Onknown	ge Areas Well y Dissolved Oxygen (mg/l) 0.38		
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)	
	23.53	9.05	-115	34	6.03	0.0	0.38	4.7	
	Cations	5/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas	
Samples Submitted to Laboratory:	Anions	6	δD a	nd $\delta^{18}O$	δ^{13} C				
				x					







Photograph Showing Location of Cottonwood Well and Spring





Cottonwood Spring Well

Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform	n Isotope San	npling and	Analysis of Gr	oundwater and Su	Irface Discharge Ar	eas		
Site Name:	Cottonwood Flume				Field Label:	N/A			
Description of Site:	Surface Water at C	ottonwood F	ume (field	data only)	Date Sampled:	7/24/2012			
	Northing:	4,032,230.5	6		Well Depth (ft):	i): N/A			
Coordinates:	Easting:	409,446.16			Screened				
	Elevation (fmsl):	3,563			Intervals (fbgs):				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)	
	19.89	7.83	-48	-16	2.63	0.0	6.22	69.3	
	Cations	s/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas	
Samples Submitted	Submitted Anions	5	δD a	and $\delta^{18}O$	δ^{13} C				
	x								





Collection of Field Data at Cottonwood Flume



View to the West of Cottonwood Flume (Note that the concrete structure is the flume used to gauge flow)





Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform	n Isotope San	npling and <i>i</i>	Analysis of Gr	oundwater and Su	urface Discharge Are	eas	
Site Name:	LADWP Cottonwoo	od Plant Well			Field Label:	Aqueduct Well		
Description of Site:	Groundwater from	LADWP Cotte	onwood Pla	ant Well	Date Sampled:	7/24/2012		
	Northing:	4,030,667.9	2		Well Depth (ft):	260		
Coordinates:	Easting:	406,827.80			Screened	206 - 260		
	Elevation (fmsl):	3,777			Intervals (fbgs):	200 - 200		
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)
	14.56	5.97	52	107	0.10	0.0	6.77	68.8
	Cations	s/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas
Samples Submitted to Laboratory:	Anions	5	δD a	nd δ ¹⁸ O	δ ¹³ C			
	X			x	x	x	x	





View of the Los Angeles Aqueduct Looking East from the Well Site





Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform	n Isotope Sam	pling and <i>i</i>	Analysis of Gr	oundwater and Su	ırface Discharge Ar	eas		
Site Name:	Ash Creek				Field Label:	Ash Crk			
Description of Site:	Surface Water from	n Ash Creek			Date Sampled:	7/24/2012			
	Northing:	4,027,549.9)		Well Depth (ft):	ell Depth (ft): N/A			
Coordinates:	Easting:	407,286.65			Screened	N/Δ			
	Elevation (fmsl):	3,872			Intervals (fbgs):				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)	
	19.59	7.96	-54	62	0.26	0.0	8.13	92	
	Cations	s/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas	
Samples Submitted to Laboratory:	Anions	S	δD a	nd δ ¹⁸ O	δ^{13} C				
	x			x					





Ash Creek Sampling Location



Collection of Field Data at Ash Creek Sampling Location





Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform	ו Isotope San	pling and	Analysis of Gr	oundwater and Su	ırface Discharge Ar	reas		
Site Name:	Rio Tinto Well				Field Label:	Rio Tinto-SW			
Description of Site:	Groundwater from	Rio Tinto We	II		Date Sampled:	7/24/2012			
	Northing:	4,022,932.0	0		Well Depth (ft):	65			
Coordinates:	Easting:	408,038.07			Screened	34 - 62			
	Elevation (fmsl):	3,616			Intervals (fbgs):	J+ - UZ			
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)	
	19.46	7.73	-40	-159	0.30	0.0	0.00	0	
	Cations/ Stable Isotopes						Tritium	Noble Gas	
Samples Submitted	Anions	5	δD a	and $\delta^{18}O$	δ^{13} C				
to Laboratory.	x			x	x	x	x		





Rio Tinto Well



View to the Southwest of Rio Tinto Well







Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform	n Isotope San	npling and <i>i</i>	Analysis of Gr	oundwater and Su	urface Discharge Are	eas	
Site Name:	Cartago Mutual Wa	ater Co. Well			Field Label:	Cartago-COM		
Description of Site:	Groundwater from	Cartago Mutu	ual Water C	o. Well	Date Sampled:	7/23/2012		
	Northing:	4020088.13			Well Depth (ft):	Unknown		
Coordinates:	Easting:	407,887.52			Screened	Unknown		
	Elevation (fmsl):	3,655			Intervals (fbgs):	OTIKIOWI		
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)
	18.00	8.33	-65.00	89.00	0.25	0.0	6.26	67.2
	Cations	s/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas
Samples Submitted to Laboratory:	Anions	8	δD a	nd δ ¹⁸ O	δ ¹³ C			
,	X			x	x	x	x	





Sample Collection at Cartago Community Well



Cartago Community Well





Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform	n Isotope San	npling and <i>i</i>	Analysis of Gr	oundwater and Su	ırface Discharge Ar	eas		
Site Name:	DWP - 9 T896				Field Label:	DWP - 9 T896			
Description of Site:	Groundwater from Monitoring Well (Ad	DWP - 9 T89 quifer 5)	6 - New OL	.GEP	Date Sampled:	7/23/2012			
	Northing:	4,041,347.6			Well Depth (ft):	j: 1,601			
Coordinates:	Easting:	412,453.5			Screened	, 1,280 - 1,360			
	Elevation (fmsl):	3,572.10			Intervals (fbgs):	1,200 - 1,300			
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)	
	19.15	9.15	-110	-349	1.46	237.0	0.00	0	
	Cation	s/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas	
Samples Submitted to Laboratory:	Anion	S	δD a	nd $\delta^{18}O$	δ^{13} C				
				x	x	x	x	x	







Noble Gas Collection at T896



Sampling at T896

Air Photo Showing Sample Site Location







OLGEP Task:	401.1.9.2 - Perforr	01.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 9 T897				Field Label:	DWP - 9 T897					
Description of Site:	Groundwater from Monitoring Well (A	DWP - 9 T897 quifer 3)	- New OL	.GEP	Date Sampled:	7/23/2012					
	Northing:	4,041,340.1	Well Depth (ft): 880								
Coordinates:	Easting:	412,453.6			Screened	780 - 860					
	Elevation (fmsl):	3,572.39			Intervals (fbgs):	780 - 880					
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)			
	19.25	9.33	-119	-412	1.77	164.0	0.00	0			
	Cations/		Stable Isot		opes	Radiocarbon	Tritium	Noble Gas			
Samples Submitted to Laboratory:	Anion	S	δD and $\delta^{18}O$		δ ¹³ C						
				x							









Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 9 T898				Field Label:	DWP - 9 T898	0WP - 9 T898			
Description of Site:	Groundwater from Monitoring Well (A	DWP - 9 T897 quifer 1)	7 - New O	LGEP	Date Sampled:	7/23/2012				
	Northing:	4,041,332.4			Well Depth (ft):): 340				
Coordinates:	Easting:	412,453.3			Screened	240 220				
	Elevation (fmsl):	3,572.22			Intervals (fbgs):					
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	19.77	9.28	-117	-349	1.35	108.0	0.00	0		
	Cation	s/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas		
Samples Submitted to Laboratory:	Anion	S	δD and $\delta^{18}O$		δ^{13} C					
				x	x	x	x	x		







Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 3 T899				Field Label:	DWP - 9 T899	Г899			
Description of Site:	Groundwater from Monitoring Well (Ad	Indwater from DWP - 3 T899 - New OLGEP itoring Well (Aquifer 5)				7/23/2012)12			
	Northing:	4,038,643.9	.9 Well Depth (ft): 1,003							
Coordinates:	Easting:	418,254.5			Screened	920 - 960				
	Elevation (fmsl):	3,572.98			Intervals (fbgs):	920 - 960				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	19.40	8.99	-101	-394	1.25	0.0	0.00	0		
	Cation	s/		Stable Isot	topes	Radiocarbon	Tritium	Noble Gas		
Samples Submitted to Laboratory:	Anion	S	δD a	and $\delta^{18}O$	δ ¹³ C					
				x	x	x	x	×		







Air Photo Showing Sample Site Location







OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 3 T901				Field Label:	DWP - 3 T901	WP - 3 T901			
Description of Site:	Groundwater from Monitoring Well (Ad	Groundwater from DWP - 3 T901 - New OLGEP Aonitoring Well (Aquifer 1)				7/23/2012	7/23/2012			
	Northing:	4,038,651.5	0		Well Depth (ft):	190				
Coordinates:	Easting:	418,265.10			Screened	150 - 170				
	Elevation (fmsl):	3,572.87			Intervals (fbgs):	150 - 170				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU) Dissolved Oxygen (mg/l)		Dissolved Oxygen (%)		
	19.78	9.26	-115	-256	3.57	406.0	0.00	0		
	Cations	s/	Stable Iso		opes	Radiocarbon	Tritium	Noble Gas		
Samples Submitted	Anions	5	δD a	nd $\delta^{18}O$	δ^{13} C					
o Laboratory:				x	x	x	x	X		







Air Photo Showing Sample Site Location







OLGEP Task:	401.1.9.2 - Perform	01.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 6 T912				Field Label:	DWP - 6 T912	DWP - 6 T912				
Description of Site:	Groundwater from Monitoring Well (A	DWP - 6 T91 quifer 5)	2 - New OL	.GEP	Date Sampled:	7/24/2012					
	Northing:	4,025,249.3			Well Depth (ft):	1,080	080				
Coordinates:	Easting:	414,248.3			Screened	1 020 - 1 060					
	Elevation (fmsl):	3,564.42			Intervals (fbgs):	1,020 - 1,000					
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)			
	20.01	9.23	-114	-316	5.58	533.0	0.00	0			
	Cation	s/		Stable Isot	topes	Radiocarbon	Tritium	Noble Gas			
Samples Submitted	Anions	5	δD a	nd $\delta^{18}O$	δ ¹³ C						
to Laboratory:				x	x	x	x				







Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform	01.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas								
Site Name:	DWP - 6 T911				Field Label:	DWP - 6 T 911	DWP - 6 T 911			
Description of Site:	Groundwater from Monitoring Well (Ad	DWP - 6 T91 quifer 5)	1 - New OL	.GEP	Date Sampled:	7/23/2012	012			
	Northing:	4,025,254.3			Well Depth (ft):	1,500				
Coordinates:	Easting:	414,252.0			Screened	1 420 - 1 460				
	Elevation (fmsl):	3,564.44			Intervals (fbgs):	1,420 - 1,400				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	19.00	8.77	-89	-303	9.82	86.7	0.00	0		
	Cation	s/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas		
Samples Submitted	Anion	5	δD and $\delta^{18}O$		δ^{13} C					
to Laboratory:				x	x	x	x	×		





DWP-6 Well Site



Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 6 T913				Field Label:	DWP - 6 T 913	DWP - 6 T 913			
Description of Site:	Groundwater from Monitoring Well (A	DWP - 6 T91 quifer 1)	3 - New OL	GEP	Date Sampled:	7/23/2012				
	Northing:	4,025,259.6			Well Depth (ft):	312				
Coordinates:	Easting:	414,255.5	; Screened 260 - 300							
	Elevation (fmsl):	3,564.51			Intervals (fbgs):	200 - 300				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	20.39	9.24	-115	-302	38.90	1.7	0.00	0		
	Cation	s/		Stable Isot	topes	Radiocarbon	Tritium	Noble Gas		
Samples Submitted	Anion	S	δD a	and $\delta^{18}O$	δ^{13} C					
to Laboratory: =				x	x	x	x			







Air Photo Showing Sample Site Location







OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 7 T908				Field Label:	DWP - 7 T908	DWP - 7 T908			
Description of Site:	Groundwater from Monitoring Well (A	DWP - 7 T90 quifer 5)	8 - New Ol	_GEP	Date Sampled:	7/23/2012				
	Northing:	4,020,292.7			Well Depth (ft):	1,470				
Coordinates:	Easting:	410,017.4			Screened	1 360 - 1 400				
	Elevation (fmsl):	3,581.90			Intervals (fbgs):	1,300 - 1,400				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	18.95	9.22	-113	-301	1.68	3.4	0.00	0		
	Cation	s/		Stable Isotopes		Radiocarbon	Tritium	Noble Gas		
Samples Submitted to Laboratory:	Anion	S	δD a	and $\delta^{18}O$	δ^{13} C					
to Laboratory:				x	x	x	x			





Installation of Monitoring Well at DWP-7





Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 7 T909				Field Label:	DWP - 7 T909				
Description of Site:	Groundwater from Monitoring Well (Ad	DWP - 7 T909 quifer 3)) - New Ol	_GEP	Date Sampled:	7/23/2012	2012			
	Northing:	4,020,298.7			Well Depth (ft):	800	0			
Coordinates:	Easting:	410,017.4			Screened	reened 740 - 780				
	Elevation (fmsl):	3,581.91			Intervals (fbgs):	740 - 780				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	19.12	9.29	-117	-315	0.75	0.0	0.00	0		
	Cation	s/		Stable Isot	topes	Radiocarbon	Tritium	Noble Gas		
Samples Submitted to Laboratory:	Anion	S	δD and $\delta^{18}O$		δ^{13} C					
				x						





Installation of Monitoring Well at DWP-7



Air Photo Showing Sample Site Location







OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	DWP - 7 T910				Field Label:	DWP - 7 T910	WP - 7 T910			
Description of Site:	Groundwater from Monitoring Well (A	DWP - 7 T91 quifer 1)	0 - New OL	.GEP	Date Sampled:	7/23/2012				
	Northing:	4,020,304.8			Well Depth (ft):	260	260			
Coordinates:	Easting:	410,018.6			Screened 200 - 240					
	Elevation (fmsl):	3,581.50			Intervals (fbgs):	200-240				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	18.76	9.03	-103	-322	0.40	0.0	0.00	0		
	Cation	s/		Stable Isotopes		Radiocarbon	Tritium	Noble Gas		
Samples Submitted to Laboratory:	Anion	S	δD a	nd $\delta^{18}O$	δ ¹³ C					
to Laboratory:				x	x	x	x	x		





Installation of Monitoring Well at DWP-7





Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform	101.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	Dolomite Well				Field Label:	Dolomite					
Description of Site:	Groundwater from	Groundwater from Dolomite Well				7/23/2012	3/2012				
	Northing: 4,046,162.64				Well Depth (ft):	127	27				
Coordinates:	Easting:	415,578.77			Screened	Inknown					
	Elevation (fmsl):	3,673			Intervals (fbgs):						
Field Parameters:	Parameters:		pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)			
	25.10	7.58	-26	-122	1.40	10.3	0.00	0			
	Cations	s/		Stable Isot	opes	Radiocarbon	Tritium	Noble Gas			
Samples Submitted to Laboratory:	Anions	8	δD a	nd δ ¹⁸ O	δ ¹³ C						
				x	x	x	x				





Dolomite Well House



Dolomite Well Inside Well House





Air Photo Showing Sample Site Location





OLGEP Task:	401.1.9.2 - Perform	01.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	AGRP-2				Field Label:	Agr-2					
Description of Site:	Groundwater from Stradling)	Groundwater from Agrarian Well No. 2 (owned by Frank Stradling)				7/23/2012					
	Northing:	4,042,586.3	5		Well Depth (ft):	140					
Coordinates:	Easting:	418,138.20		Screened 100 140							
	Elevation (fmsl):	3,604			Intervals (fbgs):						
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)			
	20.33	7.64	-28	-98	1.65	60.6	5.60	60			
	Cations	s/		Stable Isotopes		Radiocarbon	Tritium	Noble Gas			
Samples Submitted	Anions	S	δD and $\delta^{18}O$		δ^{13} C						
to Laboratory:				x	x	x	x				





View to the West of AGRP-2 Well



View to the North of AGRP-2 Well





Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	Keeler CSD Well				Field Label:	KCSD-PW				
Description of Site:	Groundwater from Keller CSD Well				Date Sampled:	7/23/2012				
Coordinates:	Northing:	4,039,398.9	3		Well Depth (ft):	125				
	Easting:	421,351.49			Screened Intervals (fbgs):	51 - 109				
	Elevation (fmsl):	3,651				51 - 103				
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	22.32	7.56	-24	22	1.59	0.0	0.00	0		
Samples Submitted to Laboratory:	Cations/ Anions		Stable Isoto		opes	Radiocarbon	Tritium	Noble Gas		
			δD and $\delta^{18}O$		δ^{13} C					
			x		x	x	x			





Sample Collection at Keeler CSD Well



Keeler CSD Well



Air Photo Showing Sample Site Location







OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	Horse Pasture Well				Field Label:	Horse Pasture				
Description of Site:	Groundwater from Horse Pasture Abandoned Flowing Well				Date Sampled:	7/23/2012				
Coordinates:	Northing:	4,039,288.6	2		Well Depth (ft):	Unknown				
	Easting:	419,995.19			Screened Intervals (fbgs):	Unknown				
	Elevation (fmsl):	3,602								
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	22.22	7.75	-35	-127	2.91	0.0	0.00	0		
Samples Submitted to Laboratory:	Cations/ Anions		Stable Isot		opes	Radiocarbon	Tritium	Noble Gas		
			δD and $\delta^{18}O$		δ^{13} C					
					x	x	x			





View to the West of Horse Pasture Well



Collection of Field Data at Horse Pasture Well





Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	S3(1) (SE Piezomter)				Field Label:	S31				
Description of Site:	Groundwater from Shallow Piezometer in SE Portion of Study Area				Date Sampled:	7/24/2012				
Coordinates:	Northing:	4,018,930.2	5		Well Depth (ft):	10				
	Easting:	411,798.53			Screened Intervals (fbgs):	9 - 10				
	Elevation (fmsl):	3,596								
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	20.00	9.43	135	-206	1.10	0.0	0.59	6.1		
Samples Submitted to Laboratory:	Cation	Stable Isot		opes	Radiocarbon	Tritium	Noble Gas			
	Anion	δD and $\delta^{18}O$		δ^{13} C						
	x x			x	x	x	x			





View to the North of S3(1) Piezometer



View to the Southeast of S3(1) Piezometer





Air Photo Showing Sample Site Location



OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	Duck 3				Field Label:	Duck-CN				
Description of Site:	Groundwater from Duck 3 Flowing Well				Date Sampled:	7/24/2012				
Coordinates:	Northing:	4,019,201.9	3		Well Depth (ft):	165				
	Easting:	411,650.52			Screened Intervals (fbgs):	40 - 140				
	Elevation (fmsl):	3,598								
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	21.08	9.22	-123	-209	0.90	0.0	1.64	18.8		
Samples Submitted to Laboratory:	Cations/		Stable Isot		opes	Radiocarbon	Tritium	Noble Gas		
	Anions		δD and $\delta^{18}O$		δ^{13} C					
		x								





Duck Pond Panorama View to the North



Air Photo Showing Sample Site Location







OLGEP Task:	401.1.9.2 - Perform Isotope Sampling and Analysis of Groundwater and Surface Discharge Areas									
Site Name:	Tubman Cement Pond				Field Label:	Tubman-CP				
Description of Site:	Spring Water from Tubman Cement Pond				Date Sampled:	7/24/2012				
Coordinates:	Northing:	4,024,798.2	5		Well Depth (ft):	N/A				
	Easting:	418,778.81			Screened Intervals (fbgs):	N/A				
	Elevation (fmsl):	3,596								
Field Parameters:	Temperature (°C)	рН	pH (mV)	Oxidation Reduction Potential (mV)	Electrical Conductivity (mS/cm)	Turbidity (NTU)	Dissolved Oxygen (mg/l)	Dissolved Oxygen (%)		
	26.43	7.94	-54	-165	8.76	893.0	2.82	36.5		
Samples Submitted to Laboratory:	Cations/		Stable Isot		opes	Radiocarbon	Tritium	Noble Gas		
	Anions		δD and $\delta^{18}O$		δ ¹³ C					
			x							





View to the North Showing Tubman Area in Background



Air Photo Showing Sample Site Location





APPENDIX B

Laboratory Analytical Data
APPENDIX B-1

Laboratory Analytical Data - General Mineral Data from LADWP Laboratory

Report No. C11913 COC 12-2032 Page 1 of 1

ENVIRONMENTAL LABORATORY DATA REPORT

Owens Lake – Dela Car Ranch Water Sample

Five water samples were collected on July 23, 2012 from the Owens Valley Dela Car Ranch.

Sample identification, tests requested, test methods, test date, and test results are listed below

All quality assurance data indicate that the results for these samples are of acceptable quality. If you have any questions, or, if further information is required, please contact Mrs. Nina Perez at (213) 367-7481 or Mr. Kevin Han at (213) 367-7267.

Sample ID	Analysis Requested	Method	Analysis Date	Results	Analysis By
LM09765	Total Dissolved Solids (TDS)	SM 2540C	7/30/2012	69 mg/L	Env Lab
	Sodium (Na)	EPA 6010B	7/31/2012	4.87 mg/L	Env Lab
1 100766	Calcium (Ca)	EPA 6010B	7/31/2012	12.8 mg/L	Env Lab
Livi09/00	Magnesium (Mg)	EPA 6010B	7/31/2012	1.49 mg/L	Env Lab
	Potassium (K)	EPA 6010B	7/31/2012	0.623 mg/L	Env Lab
I M00767	Chloride (Cl)	EPA 300.0	7/30/2012	7.22 mg/L	Env Lab
LIVI09707	Sulfate (SO_4^{-2})	EPA 300.0	7/30/2012	10.7 mg/L	Env Lab
	Alkalinity	SM 2320B	8/1/2012	42 mg/L	Env Lab
LM09768	Bicarbonate Alkalinity (HCO3)	SM 2320B	8/1/2012	42 mg/L	Env Lab
	Carbonate Alkalinity (CO ₃)	SM 2320B	8/1/2012	ND	Env Lab
LM09769	Specific Conductivity	EPA 120.1	7/30/2012	100 us/cm	Env Lab

ND: None Detected

Date Completed: 8/20/12 Work Order No.: UAE32 Job Card No.: J96745 Copies to: M.Busatto N.Liu K. Han N. Perez Filenet

Report by: Date: Checked by: Date: Approved: К. Han Date

Tests Made by: Environmental Lab

	· · . ·	Analyst(s) Assigned	Env Lab	Env Lab	Env Lab	Env Lab	Env Lab		:													Date	723/2	7125	-1/24/2	
age _1_ of _ 11 h= 2	WITTON	Test Result			-														date	fe		Time	20101	6;[]	6148	
2032 Pe	JC# U.(W T U WO#	Analysis Required	TDS	Na,Ca,Mg,K	Chloride, Sulfate	Alkalinity (CaCO ₃)/HCO3/CO3	Sp. Cond												Analvst	Approved by dat		gnature	m lie	1 2. 10	ter-	
25	() Shelf met:	Sample Matrix	water	water	water	water	L water													1		IS (- and	Y	Z	S .
110 110		ner Size	250 ml	250 ml	250 ml	250 ml	250 m			- •										297			Ka	hed by:	by	
30C #	eport C# efrig#	Contai	Ъ	- Р	- Т	- -	ъ												MOC	73:			Sampled	Relinquis	Received	
Ŭ		Preservatives	None	HNO ₃	None	None	None													Fax.						
artment of Water and Power City of Los Angeles		Sample Location and Description	Dela Car Rauch	>	\$	~	~												Marina Busatto Organization / Div.	1213 Tel 70731		Printed Name	sampled by: Sace Dorat	Relinquished by Saed Jorat	Received by D W WY	
Dep		(24 Hr) Sample Time	19:25	*	ъ.	Ĩ.	5												by.	JFB RN		Ź	reg.			ž
oratory 57	OWENS I	Sample Date	7123/12	5	×.	2	4	·				-						and the second	Requested	Address.			2-4 nis 1Day	2 Wks 4Wks	Specify	
Environmental Labo 1630 N. Main Street, Bldg Los Angeles, CA. 90012	(213) 367-7248/1399 (213) 367-7285 FAX Address Sample Location:	Chem Lab use only CHEMISTRY LOG NUMBERS (For sample duplicates use. 1 orX)	29260MJ .	2 9766	3 9767	4 9768	5 V 0769	9	7	8	5	10	11	12	13	14	15	16 Martine Martine	Date & Time		ю7 СС	DC.	12-	(fiew Lab COC Form #1 Cfiem Lab COC Form #1	Revision: 10/2/2001	

Report No. C11912 COC 12-2031 Page 1 of 1

ENVIRONMENTAL LABORATORY DATA REPORT

Owens Lake – Cartego Community Well Water Sample

Five water samples were collected on July 23, 2012 from the Owens Valley Cartego Community Well.

Sample identification, tests requested, test methods, test date, and test results are listed below

All quality assurance data indicate that the results for these samples are of acceptable quality. If you have any questions, or, if further information is required, please contact Mrs. Nina Perez at (213) 367-7481 or Mr. Kevin Han at (213) 367-7267.

Sample ID	Analysis Requested	Method	Analysis Date	Results	Analysis By
LM09760	Total Dissolved Solids (TDS)	SM 2540C	7/30/2012	138 mg/L	Env Lab
	Sodium (Na)	EPA 6010B	7/31/2012	12.1 mg/L	Env Lab
T M00761	Calcium (Ca)	EPA 6010B	7/31/2012	24.4 mg/L	Env Lab
LIV109701	Magnesium (Mg)	EPA 6010B	7/31/2012	2.07 mg/L	Env Lab
	Potassium (K)	EPA 6010B	7/31/2012	2.02 mg/L	Env Lab
L M00762	Chloride (Cl)	EPA 300.0	7/30/2012	7.37 mg/L	Env Lab
LIVI09702	Sulfate (SO_4^{-2})	EPA 300.0	7/30/2012	14.6 mg/L	Env Lab
	Alkalinity	SM 2320B	8/1/2012	80 mg/L	Env Lab
LM09763	Bicarbonate Alkalinity (HCO ₃)	SM 2320B	8/1/2012	80 mg/L	Env Lab
	Carbonate Alkalinity (CO ₃)	SM 2320B	8/1/2012	ND	Env Lab
LM09764	Specific Conductivity	EPA 120.1	7/30/2012	196 us/cm	Env Lab

ND: None Detected

Date Completed: 8/20/12 Work Order No.: UAE32 Job Card No.: J96745 Copies to: M.Busatto N.Liu K. Han N. Perez Filenet

Tests Made by: Environmental Lab Report by: Date: Checked by: Date: 612 2012 Approved: K. Han Date

	Analyst(s) Assigned	Env Lab	Env Lab	Env Lab	Env Lab	Env Lab	-														Date	7/2	1/24/V	
age_1_of UAE32	Test Result																		date e		Time	8217	817	-
- 203 1 P. JC# JAb145 WO#	Analysis Required	TDS	Na,Ca,Mg,K	Chloride, Sulfate	Alkalinity (CaCO ₃)/HCO3/CO3	Sp. Cond													Anatyst date		nature	and	- AV	0
COC #: 72 Report C# 1012 Refrig# K4 Shelf Initial of Field Personnel:	Container Sample No. Type Size Matrix	1 P 250 mL water	1 P 250 mL water	1 P 250 mL water	1 P 250 mL water	1 P 250 mL water													73297		Sampled by:	Relinquished by:	Received by	•
	Preservative	None	HNO3	None	None	None													Fax					
artment of Water and Power City of Los Angeles iin of Custody Record	Sample Location and Description	Carteyo Commenty Well	12	B		~												Marina Busatto Oreanization / Div	1213 Telt 70731	Brinted Name	Sampled by: Sae e J Jorat	Relinquished by: L J Urat	Received by D. W M	3
рер Сћа LAKE	(24 Hr) Saupie Time	1.81	L : 2	81:21 81:21	A	1018												by .	JFB RM		ES	23	120 171 G	N N
oratory	Sample Date	7123/1	4	j	2	2					-			SV1				ج: لحن Requested	Address.	Priority	2-4 Hrs 1Day	2 Wks 4Wks	Specify	
Environmental Labo 1630 N. Main Street, Bldf Los Angeles, CA. 90012 (213) 367-7248/7399 (213) 367-7285 FAX Address Sample Location:	Chern Lab use only CHEMISTRY LOG NUMBERS (For sample duplicates use. 1 or. X).	LM59760	9761	2926	1103	+0// *					0		2	0	4 00	Leten north	52 GV	Date & Time	ZIDZ CCC)C1	2- 20	31	Tevision: 10/2/2001	
Environment: 1630 N. Main Str Los Angeles, C/ (213) 367-7248/7 (213) 367-7285 F Sample Loc	Chern Lab use c CHEMISTRY LOG NU (For sample duplicates use	· LM5976	2 276		4	2	6	2	8	6	10	11	12	13	14	15	16	Date & Tin Stamp	cc)C1	2- 20	Low Lob COC Four	Tevision: 10/2/20	

Report No. C11911 COC 12-2030 Page 1 of 1

ENVIRONMENTAL LABORATORY DATA REPORT

Owens Lake – Northwest Spring Water Sample

Five water samples were collected on July 24, 2012 from the Owens Valley Northwest Spring.

Sample identification, tests requested, test methods, test date, and test results are listed below

All quality assurance data indicate that the results for these samples are of acceptable quality. If you have any questions, or, if further information is required, please contact Mrs. Nina Perez at (213) 367-7481 or Mr. Kevin Han at (213) 367-7267.

Sample ID	Analysis Requested	Method	Analysis Date	Results	Analysis By
LM09755	Total Dissolved Solids (TDS)	SM 2540C	7/30/2012	409 mg/L	Env Lab
	Sodium (Na)	EPA 6010B	7/31/2012	116 mg/L	Env Lab
I M00756	Calcium (Ca)	EPA 6010B	7/31/2012	28.2 mg/L	Env Lab
1.1109730	Magnesium (Mg)	EPA 6010B	7/31/2012	5.52 mg/L	Env Lab
	Potassium (K)	EPA 6010B	7/31/2012	5.37 mg/L	Env Lab
I M00757	Chloride (Cl)	EPA 300.0	7/30/2012	49.2 mg/L	Env Lab
L/1V109737	Sulfate (SO_4^{-2})	EPA 300.0	7/30/2012	48.1 mg/L	Env Lab
	Alkalinity	SM 2320B	8/1/2012	178 mg/L	Env Lab
LM09758	Bicarbonate Alkalinity (HCO3)	SM 2320B	8/1/2012	178 mg/L	Env Lab
	Carbonate Alkalinity (CO ₃)	SM 2320B	8/1/2012	ND	Env Lab
LM09759	Specific Conductivity	EPA 120.1	7/30/2012	672 us/cm	Env Lab

ND: None Detected

Date Completed: 8/20/12 Work Order No.: UAE32 Job Card No.: J96745 Copies to: M.Busatto N.Liu K. Han N. Perez Filenet

Tests Made by: Environmental Lab Report by: Date: Checked by Ø() Approved: K. Han Date

	Analyst(s) Assigned	Env Lab	Env Lab	Env Lab	Env Lab	Env Lab											Terlpr 7/25	1/22/2
age 1 of 1	Test Result														date te		Time 7:15 8.7(7)	\$17
ZO 30 P. P. July June P. July June P. July July Wolf Wolf Wolf Wolf Wolf Wolf P. No. of Field Test:	Analysis Required	TDS	Na,Ca,Mg,K	Chloride, Sulfate	Alkalinity (CaCO ₃)/HCO3/CO3	Sp. Cond	-								Analyst		matyre pro	
12 - 11 MII Shelf	Sample Matrix	ıL water	1L water	ıL water	nL water	ון water											Blig	A
C# 10 C# 10 Field Person	itainer _{Ype} Size	P 250 m	P 250 m	P 250 n	P 250 n	P 250 n				 					73297		at hu Cless	
COC Report Refrig	No. Cor	-		~		-		 							MQQ		Sample Reling	
	Preservatives	None	HNO3	None	None	None									Чях			
urtment of Water and Power City of Los Angeles in of Custody Record #2	Sample Location and Description	Northwest Spring				-		· · · · · · · · · · · · · · · · · · ·							Marina Busatto Organization / Div. 1213 Telk 70731		Printed Name Sampled by: Saeed Jorat Relinquished by: Jorat Received by	Cum"
^{Dep} Сha	24 Hr) Sample Time	1.0	•+					 							y . JFB RM		至至男方	医が
oratory g ⁷ OWENS L	Sample Date	7/24/42						 				CO			Requested b Address.		2.4 Hrs 2.4 Hrs 1Day 2 Wks 4 Wks Specify	
Environmental Labo 1630 N. Main Street, Bld Los Angeles, CA. 90012 (213) 367-7248/7399 (213) 367-7285 FAX Address Sample Location:	Chern Lab use only CHEMISTRY LOG NUMBERS (For sample duplicates use .1 or .X)	1M 69755	9756	L526	9758	V 9759			10	12	13	14 0	15 🖓 👘	16 A 45	Date & Time Stamp Stamp Stamp	ና ፖ DC1	13 How Tab COC Form #	reason:

Report No. C11910 COC 12-2029 Page 1 of 1

ENVIRONMENTAL LABORATORY DATA REPORT

Owens Lake – Cottonwood Treatment Plant Well Water Sample

Five water samples were collected on July 24, 2012 from the Owens Valley Cottonwood Treatment Plant Well.

Sample identification, tests requested, test methods, test date, and test results are listed below

All quality assurance data indicate that the results for these samples are of acceptable quality. If you have any questions, or, if further information is required, please contact Mrs. Nina Perez at (213) 367-7481 or Mr. Kevin Han at (213) 367-7267.

Sample ID	Analysis Requested	Method	Analysis Date	Results	Analysis By
LM09750	Total Dissolved Solids (TDS)	SM 2540C	7/30/2012	56 mg/L	Env Lab
	Sodium (Na)	EPA 6010B	7/31/2012	4.16 mg/L	Env Lab
I M00751	Calcium (Ca)	EPA 6010B	7/31/2012	7.19 mg/L	Env Lab
LIVI09/31	Magnesium (Mg)	EPA 6010B	7/31/2012	1.75 mg/L	Env Lab
	Potassium (K)	EPA 6010B	7/31/2012	1.43 mg/L	Env Lab
I M00750	Chloride (Cl)	EPA 300.0	7/30/2012	7.07 mg/L	Env Lab
LIVI09/32	Sulfate (SO_4^{-2})	EPA 300.0	7/30/2012	9.87 mg/L	Env Lab
	Alkalinity	SM 2320B	8/1/2012	30 mg/L	Env Lab
LM09753	Bicarbonate Alkalinity (HCO ₃)	SM 2320B	8/1/2012	30 mg/L	Env Lab
	Carbonate Alkalinity (CO ₃)	SM 2320B	8/1/2012	ND	Env Lab
LM09754	Specific Conductivity	EPA 120.1	7/30/2012	73 us/cm	Env Lab

ND: None Detected

Date Completed: 8/20/12 Work Order No.: UAE32 Job Card No.: J96745 Copies to: M.Busatto N.Liu K. Han N. Perez Filenet

Tests Made by: Environmental Lab Report by: NP Date: 8/20/12 Checked by: Date: 6/20/12 Approved: K. Han Date

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oratory	Sample Date	21/26/1	1 N	, N	- A - A 	-								-gyr			a Requested	Priority	2-4 Hrs 1Day 2 Wks 4Wks Specify	
Environmental Labo 1630 N. Main Street, Bld Los Angeles, CA. 90012 (213) 367-7248/7399 (213) 367-7285 FAX Address Sample Location:	Chem Lab use only CHEMISTRY LOG NUMBERS (For sample duplicates use .1 or .X)	1 LM 69750	2 4751 3 67 m	4753	5 J 9754	Q	<u> </u>	8	6	10	11	12	13	14	15 Š	16 • • •	2 International Constraints	تیں 20 20	nd 10/2/2001	

Report No. C11914 COC 12-2033 Page 1 of 1

ENVIRONMENTAL LABORATORY DATA REPORT

Owens Lake – Cottonwood Spring Northwest Water Sample

Five water samples were collected on July 24, 2012 from the Owens Valley Cottonwood Spring Northwest.

Sample identification, tests requested, test methods, test date, and test results are listed below

All quality assurance data indicate that the results for these samples are of acceptable quality. If you have any questions, or, if further information is required, please contact Mrs. Nina Perez at (213) 367-7481 or Mr. Kevin Han at (213) 367-7267.

Sample ID	Analysis Requested	Method	Analysis Date	Results	Analysis By
LM09770	Total Dissolved Solids (TDS)	SM 2540C	7/30/2012	230 mg/L	Env Lab
	Sodium (Na)	EPA 6010B	7/31/2012	29.0 mg/L	Env Lab
I M00771	Calcium (Ca)	EPA 6010B	7/31/2012	26.5 mg/L	Env Lab
LJV109771	Magnesium (Mg)	EPA 6010B	7/31/2012	7.21 mg/L	Env Lab
	Potassium (K)	EPA 6010B	7/31/2012	3.55 mg/L	Env Lab
I M00772	Chloride (Cl)	EPA 300.0	7/30/2012	42.3 mg/L	Env Lab
LIV109772	Sulfate (SO ₄ ⁻²)	EPA 300.0	7/30/2012	23.9 mg/L	Env Lab
	Alkalinity	SM 2320B	8/1/2012	70 mg/L	Env Lab
LM09773	Bicarbonate Alkalinity (HCO ₃)	SM 2320B	8/1/2012	70 mg/L	Env Lab
	Carbonate Alkalinity (CO ₃)	SM 2320B	8/1/2012	ND	Env Lab
LM09774	Specific Conductivity	EPA 120.1	7/30/2012	347 us/cm	Env Lab

ND: None Detected

Date Completed: 8/20/12 Tests Made by: Environmental Lab Work Order No.: UAE32 Job Card No.: J96745 Report by: Date: Copies to: M.Busatto Checked by: Date: N.Liu K. Han Approved: N. Perez K. Han Date Acting Manager of Environmental Laboratory Filenet

	Analyst(s) Assigned	Env Lab	Env Lab	Env Lab	Env Lab	ENV LaD														Date	212412	7/25	2/24/2	81/52/-1
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oratory g ^{g 7} OWENS I	Sample Date	21/22/12	2 2		4									to the second		Landara Sectors Sectors Sectors	Requested 1	Address.	<₿riority	CE 2-4 Hrs	1Day 2 Wks	4Wks Specify		
Environmental Labu 1630 N. Main Street, Bldi Los Angeles, CA. 90012 (213) 367-7248/7399 (213) 367-7285 FAX Address Sample Location:	Chem Lab use only CHEMISTRY LOG NUMBERS (For sample duplicates use 1 or X)	02290M1	1116 2	4 9773	5 V 9774	0	2	8	6	10	11	12	13	14	15 600		Stamp Stamp	Ci Ci	کتر. کرر	12-	20:	hem Lab COC Form #1	Revision: 10/2/2001	

Report No. C11915 COC 12-2034 Page 1 of 1

ENVIRONMENTAL LABORATORY DATA REPORT

Owens Lake – S(3)1 Pizometer Water Sample

Five water samples were collected on July 24, 2012 from the Owens Valley S(3)1 Pizometer.

Sample identification, tests requested, test methods, test date, and test results are listed below

All quality assurance data indicate that the results for these samples are of acceptable quality. If you have any questions, or, if further information is required, please contact Mrs. Nina Perez at (213) 367-7481 or Mr. Kevin Han at (213) 367-7267.

Sample ID	Analysis Requested	Method	Analysis Date	Results	Analysis By
LM09775	Total Dissolved Solids (TDS)	SM 2540C	7/30/2012	555 mg/L	Env Lab
	Sodium (Na)	EPA 6010B	7/31/2012	179 mg/L	Env Lab
I M00776	Calcium (Ca)	EPA 6010B	7/31/2012	1.71 mg/L	Env Lab
LIV109770	Magnesium (Mg)	EPA 6010B	7/31/2012	0.051 mg/L	Env Lab
	Potassium (K)	EPA 6010B	7/31/2012	15.8 mg/L	Env Lab
I M00777	Chloride (Cl)	EPA 300.0	7/30/2012	92.8 mg/L	Env Lab
LIVIU	Sulfate (SO ₄ ⁻²)	EPA 300.0	7/30/2012	79.5 mg/L	Env Lab
	Alkalinity	SM 2320B	8/1/2012	150 mg/L	Env Lab
LM09778	Bicarbonate Alkalinity (HCO3)	SM 2320B	8/1/2012	90 mg/L	Env Lab
	Carbonate Alkalinity (CO ₃)	SM 2320B	8/1/2012	60 mg/L	Env Lab
LM09779	Specific Conductivity	EPA 120.1	7/30/2012	886 us/cm	Env Lab

ND: None Detected

Date Completed: 8/20/12 Tests Made by: Environmental Lab Work Order No.: UAE32 Job Card No.: J96745 Date: 8/20/12 Report by: NP Checked by: SL Date: k Copies to: M.Busatto N.Liu K. Han Approved: N. Perez Date Han Filenet Acting Manager of Environmental Laboratory

	Analyst(s) Assigned Env Lab Env Lab	Env Lab Env Lab Env Lab Date
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oratory ¹⁸ 7 2 S	Sample Date	Specify Specify
Environmental Lab 1630 N. Main Street, Bld Los Angeles, CA. 9001: (213) 367-7248/7399 (213) 367-7285 FAX (213) 367-7285 FAX Address Sample Location:	Chem Lab use only CHEMISTRY LOG NUMBERS For sample duplicates use 1 or XI LM04775 2 9776 3 9777	6 6 7

Report No. C11916 COC 12-2035 Page 1 of 1

ENVIRONMENTAL LABORATORY DATA REPORT

Owens Lake – Ash Creek Water Sample

Five water samples were collected on July 24, 2012 from the Owens Valley Ash Creek.

Sample identification, tests requested, test methods, test date, and test results are listed below

All quality assurance data indicate that the results for these samples are of acceptable quality. If you have any questions, or, if further information is required, please contact Mrs. Nina Perez at (213) 367-7481 or Mr. Kevin Han at (213) 367-7267.

Sample ID	Analysis Requested	Method	Analysis Date	Results	Analysis By
LM09780	Total Dissolved Solids (TDS)	SM 2540C	7/30/2012	122 mg/L	Env Lab
	Sodium (Na)	EPA 6010B	7/31/2012	7.95 mg/L	Env Lab
T M00791	Calcium (Ca)	EPA 6010B	7/31/2012	23.7 mg/L	Env Lab
LW09/81	Magnesium (Mg)	EPA 6010B	7/31/2012	3.67 mg/L	Env Lab
	Potassium (K)	EPA 6010B	7/31/2012	1.87 mg/L	Env Lab
T M00792	Chloride (Cl)	EPA 300.0	7/30/2012	7.37 mg/L	Env Lab
LIVI09/82	Sulfate (SO_4^{-2})	EPA 300.0	7/30/2012	29.0 mg/L	Env Lab
	Alkalinity	SM 2320B	8/1/2012	60 mg/L	Env Lab
LM09783	Bicarbonate Alkalinity (HCO ₃)	SM 2320B	8/1/2012	60 mg/L	Env Lab
	Carbonate Alkalinity (CO ₃)	SM 2320B	8/1/2012	ND	Env Lab
LM09784	Specific Conductivity	EPA 120.1	7/30/2012	198 mg/L	Env Lab

ND: None Detected

Date Completed: 8/20/12 Work Order No.: UAE32 Job Card No.: J96745 Copies to: M.Busatto N.Liu K. Han N. Perez Filenet

Tests Made by: Environmental Lab Report by: Date: Checked by: Date: Approved: K. Han

2035 Page_1_of_1_ UNDTUS_wo#_UAE32- S\$_Bin# No. of Field Test	Analysis Test Analysi(s) Required Result Assigned	TDS Env Lab	Na,Ca,Mg,K Env Lab	Chloride, Sulfate Env Lab	alinity (CaCO ₃)/HCO3/CO3	Sp. Cond Env Lab												alyst date date	ITTIME Date	20) 8:17 7/25	8:17 7/22/-
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Report No. C11917 COC 12-2036 Page 1 of 1

ENVIRONMENTAL LABORATORY DATA REPORT

Owens Lake – Rio Tinto Supply Well Water Sample

Five water samples were collected on July 24, 2012 from the Owens Valley Rio Tinto Supply Well.

Sample identification, tests requested, test methods, test date, and test results are listed below

All quality assurance data indicate that the results for these samples are of acceptable quality. If you have any questions, or, if further information is required, please contact Mrs. Nina Perez at (213) 367-7481 or Mr. Kevin Han at (213) 367-7267.

Sample ID	Analysis Requested	Method	Analysis Date	Results	Analysis By
LM09785	Total Dissolved Solids (TDS)	SM 2540C	7/30/2012	158 mg/L	Env Lab
	Sodium (Na)	EPA 6010B	7/31/2012	25.9 mg/L	Env Lab
T. N.60079.C	Calcium (Ca)	EPA 6010B	7/31/2012	17.0 mg/L	Env Lab
LIM09/86	Magnesium (Mg)	EPA 6010B	7/31/2012	2.14 mg/L	Env Lab
	Potassium (K)	EPA 6010B	7/31/2012	3.28 mg/L	Env Lab
1.) (00797	Chloride (Cl)	EPA 300.0	7/30/2012	7.87 mg/L	Env Lab
LIV109/8/	Sulfate (SO ₄ ⁻²)	EPA 300.0	7/30/2012	15.7 mg/L	Env Lab
	Alkalinity	SM 2320B	8/1/2012	94 mg/L	Env Lab
LM09788	Bicarbonate Alkalinity (HCO ₃)	SM 2320B	8/1/2012	94 mg/L	Env Lab
	Carbonate Alkalinity (CO ₃)	SM 2320B	8/1/2012	ND	Env Lab
LM09789	Specific Conductivity	EPA 120.1	7/30/2012	226 mg/L	Env Lab

ND: None Detected

Date Completed: 8/20/12 Work Order No.: UAE32 Job Card No.: J96745 Copies to: M.Busatto N.Liu K. Han N. Perez Filenet

Tests Made by: Environmental Lab Date: Report by: NP -Date: 812 Checked by: Approved: K. Han Date

Environmel 1630 N. Main Los Angeles, (213) 367-724 (213) 37-724 (213)	ntal Labo Street, Bldi CA. 90012 8/7399 5 FAX Address only Se only Se	Clear AB Clear Covens	Dep Dep Cha	artment of Water and Power City of Los Angeles in of Custody Record # 8 # 8 # 8 # 8 Sample Location and Description Rie Tinto Supply Well	Preservatives None None None	COC Con Initial of Initia of Initial of Initial of Initial of Initial of Init	# # : C# #// ainer ainer 250 n 250	Shelf Sample	- ZO36 P IC# MOTHT WO# S Bin#. WO# Analysis Required TDS Analysis Required Analysis Required Analysis Required Analysis Required TDS Na, of Field Test: Sp. Cond	LAEST OF Lof Lest Result	Analyst(s Env Lab Env Lab Env Lab Env Lab	
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APPENDIX B-2

Laboratory Analytical Data - Isotope Data from Isotech Laboratories

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MULLIA MULLIA<	Project Contact: Name: Victor Harris, MWH Americas, Inc. Phone: 626-568-6024 Email: Victor E Harris@us mvhelobal.com				Å			Namo Phon Email: le	Lab Cont. c: Christy hc: 877-63 procr@isot	act / Legner (2:4190 techlabs.c	Ш		
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Horse Pasture 33 $723,70,1$ x	KCSD-PW	ŝ			7/23/2012		×	×	x	×	×		
531 531 724,7012 x x x DuckCN 1 1 724,7012 x x DuckCN 1 1 724,7012 x x DuckCN 1 1 1 1 x DuckCN 1 1 1 1 1 DuckCN 1	Horse Pasture	3			7/23/2012		x	x	X	x	x		
$\begin{array}{ c c c c c c c c c c c c c c c c c c c$	531	ε			7/24/2012		×	×	x	×	×		
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Relinquished by: M. Lee Mauli SCOM Date/Time Received by: Received by: Date/Time	Tubman-CP	1			7/24/2012		x	×					
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Isotech Water Data Job 18823

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Isotech	Sample	Sample	Field	Analysis	$\delta D H_2 O$	$\delta^{18}OH_2O$	Tritium	Std. Dev.	δ^{13} C DIC	¹⁴ C DIC	Std. Dev.
Lab No.	Name	Date	Name	Date	‰	‰	TU		‰	pMC	
258207	Carroll Crk	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-120.5	-15.94					
258208	NW Seep	7/24/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-121.4	-15.80	< 1.00		-8.9	57.5	0.2
258209	Bartlett-CW	7/24/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-123.3	-16.16	< 1.00		-9.3	52.8	0.2
258210	Cottonwood Spr	7/24/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-120.4	-15.47	7.46	0.20	-15.7	88.9	0.3
258211	Cottonwood-Well	7/24/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-103.5	-12.01					
258212	Aqueduct Well	7/24/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-119.3	-15.40	3.82	0.17	-15.1	106.5	0.4
258213	Ash Crk	7/24/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-111.2	-14.56					
258214	Rio Tinto-SW	7/24/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-114.5	-14.88	< 1.00		-9.8	56.1	0.2
258215	Cartago-COM	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-112.1	-14.71	1.76	0.15	-10.4	61.3	0.3
258216	DWP-9 / T 896	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-129.5	-16.55	< 1.00		21.7	2.9	0.1
258217	DWP-9 / T 897	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-131.8	-16.96					
258218	DWP-9 / T 898	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-129.4	-16.74	< 1.00		5.0	7.2	0.1
258219	DWP-3 / T 899	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-123.3	-15.68	< 1.00		-1.5	3.8	0.1
258220	DWP-3 / T 901	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-125.8	-16.11	< 1.00		0.7	1.8	0.1
258221	DWP-6 / T 912	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-132.7	-16.59	< 1.00		-0.2	1.7	0.1
258222	DWP-6 / T 911	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-114.2	-13.25	< 1.00		-7.7	1.2	0.1
258223	DWP-6 / T 913	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-88.0	-8.28	< 1.00		0.9	0.5	0.1
258224	DWP-7 / T 908	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-108.9	-14.34	< 1.00		-8.1	4.4	0.1
258225	DWP-7 / T 909	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-109.0	-14.38					
258226	DWP-7 / T 910	7/23/2012	OLGEP TASK: 401.1.9.2	8/5/2012	-110.1	-14.62	< 1.00		-6.0	36.7	0.2
258227	Dolomite	7/23/2012	OLGEP TASK: 401.1.9.2	8/6/2012	-127.6	-16.29	< 1.00		-7.6	5.6	0.1
258228	Agr-2	7/23/2012	OLGEP TASK: 401.1.9.2	8/6/2012	-125.0	-16.13	< 1.00		-5.8	3.6	0.1
258229	KCSD-PW	7/23/2012	OLGEP TASK: 401.1.9.2	8/6/2012	-124.5	-16.09	< 1.00		-5.5	4.4	0.1
258230	Horse Pasture	7/23/2012	OLGEP TASK: 401.1.9.2	8/6/2012	-130.9	-16.70	< 1.00		-5.1	1.9	0.1
258231	S31	7/24/2012	OLGEP TASK: 401.1.9.2	8/6/2012	-110.2	-14.28	< 1.00		-3.7	25.5	0.2
258232	Duck-CN	7/24/2012	OLGEP TASK: 401.1.9.2	8/6/2012	-111.7	-14.82					
258233	Tubman-CP	7/24/2012	OLGEP TASK: 401.1.9.2	8/6/2012	-128.8	-15.66					



Water Analysis

Lab Number:	258207		Job Number:	18823
Submitter Sample Name:	Carroll Crk			
Submitter Sample ID:				
Submitter Job #:				
Company:	MWH Labo	ratories		
Field or Site:	OLGEP TA	SK: 401.1.9.2		
Location:				
Depth/Formation:				
Container Type:	Amber Bottl	е		
Sample Collected:	7/23/2012	Res	ults Reported:	8/31/2012
δD of water .		-120.5 ‰ rel	ative to VSMOW	
$\delta^{18}O$ of water .		-15.94 ‰ rel	ative to VSMOW	
Tritium content of water -		na		
$\delta^{13}C$ of DIC .		na		
¹⁴ C content of DIC		na		
$\delta^{15}N$ of nitrate .		na		
$\delta^{18}O$ of nitrate .		na		
$\delta^{34}S$ of sulfate .		na		
$\delta^{18}O$ of sulfate $$-$$		na		



Water Analysis

Lab Number:	258208			Job Number:	18823
Submitter Sample Name:	NW Seep				
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	atories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/24/2012		Results F	Reported:	8/31/2012
δD of water		-121.4 %	‰ relative	to VSMOW	
δ^{18} O of water		-15.80 %	‰ relative	to VSMOW	
Tritium content of water-		< 1.00	TU		
δ^{13} C of DIC		-8.9 ‰ r	elative to	VPDB	
¹⁴ C content of DIC		57.5 ±	0.2 perce	ent modern ca	rbon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258209			Job Number:	18823
Submitter Sample Name:	Bartlett-CW				
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/24/2012		Results	Reported:	8/31/2012
δD of water		-123.3 %	‰ relative	e to VSMOW	
δ^{18} O of water		-16.16 %	la relative	e to VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-9.3 ‰ r	elative to	VPDB	
¹⁴ C content of DIC		52.8 ±	0.2 perc	ent modern ca	rbon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258210			Job Number:	18823
Submitter Sample Name:	Cottonwood	Spr			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	atories			
Field or Site:	OLGEP TAS	SK: 401.	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/24/2012		Results	Reported:	8/31/2012
δD of water	·	-120.4 🤅	‰ relative	e to VSMOW	
δ^{18} O of water	·	-15.47 %	‰ relative	e to VSMOW	
Tritium content of water -		7.46 ±	0.20 TU		
δ^{13} C of DIC		-15.7 %	o relative	to VPDB	
¹⁴ C content of DIC		88.9 ±	0.3 perc	ent modern ca	rbon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate	·	na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258211			Job Number:	18823
Submitter Sample Name	: Cottonwood	d-Well			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labo	ratories			
Field or Site:	OLGEP TA	SK: 401.	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bott	le			
Sample Collected:	7/24/2012		Results	Reported:	8/31/2012
δD of water		-103.5 %	‰ relative	to VSMOW	
$\delta^{18}O$ of water		-12.01 %	‰ relative	to VSMOW	
Tritium content of water	·	na			
$\delta^{13}C$ of DIC		na			
¹⁴ C content of DIC		na			
$\delta^{15}N$ of nitrate		na			
$\delta^{18}O$ of nitrate		na			
$\delta^{34}S$ of sulfate		na			
$\delta^{18}O$ of sulfate		na			



Water Analysis

Lab Number:	258212			Job Number:	18823
Submitter Sample Name:	Aqueduct W	/ell			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401.1	.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/24/2012		Results I	Reported:	8/31/2012
δD of water		-119.3 %	 relative 	to VSMOW	
δ^{18} O of water		-15.40 %	 relative 	to VSMOW	
Tritium content of water		3.82 ±0	0.17 TU		
δ^{13} C of DIC		-15.1 ‰	relative t	o VPDB	
¹⁴ C content of DIC		106.5 ±	0.4 perc	ent modern ca	arbon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258213		Job Number:	18823
Submitter Sample Name:	Ash Crk			
Submitter Sample ID:				
Submitter Job #:				
Company:	MWH Labo	ratories		
Field or Site:	OLGEP TA	SK: 401.1.9.2		
Location:				
Depth/Formation:				
Container Type:	Amber Bottl	le		
Sample Collected:	7/24/2012	Res	ults Reported:	8/31/2012
δD of water .		-111.2 ‰ rel	ative to VSMOW	
δ^{18} O of water		-14.56 ‰ rel	ative to VSMOW	
Tritium content of water -		na		
$\delta^{13}C$ of DIC .		na		
¹⁴ C content of DIC		na		
$\delta^{15}N$ of nitrate .		na		
δ^{18} O of nitrate		na		
$\delta^{34}S$ of sulfate		na		
$\delta^{18}O$ of sulfate .		na		



Water Analysis

Lab Number:	258214			Job Number:	18823
Submitter Sample Name:	Rio Tinto-SV	N			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	atories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/24/2012		Results	Reported:	8/31/2012
δD of water	·	-114.5 %	‰ relative	e to VSMOW	
δ^{18} O of water	·	-14.88 %	‰ relative	e to VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-9.8 ‰ ı	elative to	VPDB	
¹⁴ C content of DIC		56.1 ±	0.2 perc	ent modern ca	rbon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate	·	na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258215			Job Number:	18823
Submitter Sample Name:	Cartago-CC	M			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labo	ratories			
Field or Site:	OLGEP TAS	SK: 401.	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/23/2012		Results	Reported:	8/31/2012
δD of water		-112.1 %	‰ relative	e to VSMOW	
δ^{18} O of water		-14.71 %	‰ relative	e to VSMOW	
Tritium content of water		1.76 ±	0.15 TU		
δ^{13} C of DIC		-10.4 %	relative	to VPDB	
¹⁴ C content of DIC		61.3 ±	0.3 perc	cent modern ca	rbon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258216		·	Job Number:	18823
Submitter Sample Name:	DWP-9 / T 8	396			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/23/2012		Results F	Reported:	8/31/2012
δD of water		-129.5 %	‰ relative	to VSMOW	
δ^{18} O of water		-16.55 %	‰ relative	to VSMOW	
Tritium content of water -		< 1.00	TU		
δ^{13} C of DIC		21.7 ‰	relative to	VPDB	
¹⁴ C content of DIC		2.9 ±	0.1 perce	nt modern car	bon
δ^{15} N of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258217		Job Number:	18823
Submitter Sample Name:	DWP-9 / T 8	897		
Submitter Sample ID:				
Submitter Job #:				
Company:	MWH Labo	ratories		
Field or Site:	OLGEP TA	SK: 401.1.9.2		
Location:				
Depth/Formation:				
Container Type:	Amber Bottl	e		
Sample Collected:	7/23/2012	Res	ults Reported:	8/31/2012
δD of water .		-131.8 ‰ rel	ative to VSMOW	
δ^{18} O of water		-16.96 ‰ rel	ative to VSMOW	
Tritium content of water -		na		
$\delta^{13}C$ of DIC		na		
¹⁴ C content of DIC		na		
$\delta^{15}N$ of nitrate		na		
δ^{18} O of nitrate		na		
$\delta^{34}S$ of sulfate		na		
δ^{18} O of sulfate		na		



Water Analysis

Lab Number:	258218		Job	Number:	18823
Submitter Sample Name:	DWP-9 / T 8	398			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/23/2012		Results Repo	orted:	8/31/2012
δD of water		-129.4 %	‰ relative to V	SMOW	
δ^{18} O of water		-16.74 %	‰ relative to V	SMOW	
Tritium content of water-		< 1.00	TU		
δ^{13} C of DIC		5.0 ‰ re	elative to VPD	В	
¹⁴ C content of DIC		7.2 ±	0.1 percent m	odern car	bon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
δ ³⁴ S of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258219		· ·	Job Number:	18823
Submitter Sample Name:	DWP-3 / T 8	399			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401.	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/23/2012		Results F	Reported:	8/31/2012
δD of water		-123.3 %	‰ relative	to VSMOW	
δ^{18} O of water		-15.68 %	‰ relative	to VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-1.5 ‰ เ	relative to	VPDB	
¹⁴ C content of DIC		3.8 ±	0.1 percer	nt modern car	bon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258220		J	ob Number:	18823
Submitter Sample Name:	DWP-3 / T 9	901			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401.1	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/23/2012		Results Re	eported:	8/31/2012
δD of water		-125.8 %	lo relative to	o VSMOW	
δ^{18} O of water		-16.11 %	lo relative to	o VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		0.7 ‰ re	elative to V	PDB	
¹⁴ C content of DIC		1.8 ±	0.1 percen	t modern car	bon
δ^{15} N of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258221		·	Job Number:	18823
Submitter Sample Name:	DWP-6 / T 9	912			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	atories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/23/2012		Results F	Reported:	8/31/2012
δD of water	·	-132.7 %	‰ relative	to VSMOW	
δ^{18} O of water	·	-16.59 %	‰ relative	to VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-0.2 ‰ r	relative to	VPDB	
¹⁴ C content of DIC		1.7 ±	0.1 perce	nt modern car	bon
δ^{15} N of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate	·	na			



Water Analysis

258222			Job Number:	18823
DWP-6 / T 9	911			
MWH Labo	ratories			
OLGEP TAS	SK: 401. ⁻	1.9.2		
Amber Bottl	е			
7/23/2012		Results F	Reported:	8/31/2012
	-114.2 %	‰ relative	to VSMOW	
	-13.25 %	‰ relative	to VSMOW	
	< 1.00	TU		
	-7.7 ‰ I	relative to	VPDB	
	1.2 ±	0.1 perce	nt modern car	bon
	na			
	258222 DWP-6 / T S MWH Labor OLGEP TAS Amber Bottl 7/23/2012	258222 DWP-6 / T 911 MWH Laboratories OLGEP TASK: 401. Amber Bottle 7/23/2012 	258222 A DWP-6 / T 911 MWH Laboratories OLGEP TASK: 401.1.9.2 Amber Bottle 7/23/2012 Results F -114.2 % relative -13.25 % relative -1.2 ± 0.1 perce na na na na na na na na	258222 Job Number: DWP-6 / T 911 MWH Laboratories OLGEP TASK: 401.1.9.2 Amber Bottle 7/23/2012 Results Reported: -114.2 % relative to VSMOW -13.25 % relative to VSMOW -1.2 ± 0.1 percent modern car na na na na


Water Analysis

Lab Number:	258223			Job Number:	18823
Submitter Sample Name:	DWP-6 / T 9	913			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	atories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/23/2012		Results F	Reported:	8/31/2012
δD of water		-88.0 ‰	relative to	o VSMOW	
δ^{18} O of water		-8.28 ‰	relative to	o VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC	·	0.9 ‰ re	elative to '	VPDB	
¹⁴ C content of DIC		0.5 ±	0.1 perce	ent modern car	bon
δ^{15} N of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate	·	na			



Water Analysis

Lab Number:	258224		Jo	ob Number:	18823
Submitter Sample Name:	DWP-7 / T 9	908			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/23/2012		Results Re	eported:	8/31/2012
δD of water	·	-108.9 %	‰ relative to	SMOW	
δ^{18} O of water	·	-14.34 %	‰ relative to	SMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-8.1 ‰ r	relative to V	'PDB	
¹⁴ C content of DIC		4.4 ±	0.1 percent	t modern car	bon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate	·	na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258225		Job Number:	18823
Submitter Sample Name:	DWP-7 / T 9	909		
Submitter Sample ID:				
Submitter Job #:				
Company:	MWH Labo	ratories		
Field or Site:	OLGEP TAS	SK: 401.1.9.2		
Location:				
Depth/Formation:				
Container Type:	Amber Bottl	е		
Sample Collected:	7/23/2012	Result	s Reported:	8/31/2012
δD of water .		-109.0 ‰ relativ	ve to VSMOW	
δ^{18} O of water .		-14.38 ‰ relativ	ve to VSMOW	
Tritium content of water -		na		
$\delta^{13}C$ of DIC .		na		
¹⁴ C content of DIC		na		
$\delta^{15}N$ of nitrate .		na		
$\delta^{18}O$ of nitrate .		na		
$\delta^{34}S$ of sulfate .		na		
$\delta^{18}O$ of sulfate .		na		



Water Analysis

Lab Number:	258226			Job Number:	18823
Submitter Sample Name:	DWP-7 / T 9	910			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	atories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/23/2012		Results I	Reported:	8/31/2012
δD of water		-110.1 %	‰ relative	to VSMOW	
δ^{18} O of water		-14.62 %	lo relative	to VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-6.0 ‰ r	elative to	VPDB	
¹⁴ C content of DIC		36.7 ±	0.2 perc	ent modern ca	rbon
$\delta^{15}N$ of nitrate		na			
δ ¹⁸ O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate	·	na			



Water Analysis

/31/2012



Water Analysis

Lab Number:	258228		Jo	ob Number:	18823
Submitter Sample Name:	Agr-2				
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labo	ratories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/23/2012		Results Re	eported:	8/31/2012
δD of water		-125.0 %	‰ relative to	SMOW	
δ^{18} O of water		-16.13 %	‰ relative to	SMOW	
Tritium content of water -		< 1.00	TU		
δ ¹³ C of DIC		-5.8 ‰ ı	relative to V	'PDB	
¹⁴ C content of DIC		3.6 ±	0.1 percen	t modern car	bon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258229		· ·	Job Number:	18823
Submitter Sample Name:	KCSD-PW				
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	atories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/23/2012		Results F	Reported:	8/31/2012
δD of water	·	-124.5 %	‰ relative	to VSMOW	
δ^{18} O of water	·	-16.09 %	‰ relative	to VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-5.5 ‰ r	relative to	VPDB	
¹⁴ C content of DIC		4.4 ±	0.1 percer	nt modern car	bon
δ^{15} N of nitrate		na			
δ^{18} O of nitrate	·	na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258230		Job N	lumber:	18823
Submitter Sample Name:	Horse Pastu	ıre			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401.	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	е			
Sample Collected:	7/23/2012		Results Repor	ted:	8/31/2012
δD of water		-130.9 %	‰ relative to VS	SMOW	
δ^{18} O of water		-16.70 %	‰ relative to VS	SMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-5.1 ‰ I	elative to VPD	В	
¹⁴ C content of DIC		1.9 ±	0.1 percent mo	odern car	bon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258231			Job Number:	18823
Submitter Sample Name:	S31				
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labor	ratories			
Field or Site:	OLGEP TAS	SK: 401. ⁻	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bottl	e			
Sample Collected:	7/24/2012		Results	Reported:	8/31/2012
δD of water		-110.2 %	la relative	e to VSMOW	
δ^{18} O of water		-14.28 %	‰ relative	e to VSMOW	
Tritium content of water		< 1.00	TU		
δ^{13} C of DIC		-3.7 ‰ r	elative to	VPDB	
¹⁴ C content of DIC		25.5 ±	0.2 perc	ent modern ca	rbon
$\delta^{15}N$ of nitrate		na			
δ^{18} O of nitrate		na			
$\delta^{34}S$ of sulfate		na			
δ^{18} O of sulfate		na			



Water Analysis

Lab Number:	258232		Job Number:	18823
Submitter Sample Name:	Duck-CN			
Submitter Sample ID:				
Submitter Job #:				
Company:	MWH Labo	ratories		
Field or Site:	OLGEP TA	SK: 401.1.9.2		
Location:				
Depth/Formation:				
Container Type:	Amber Bottl	е		
Sample Collected:	7/24/2012	Resu	ults Reported:	8/31/2012
δD of water .		-111.7 ‰ rela	ative to VSMOW	
$\delta^{18}O$ of water .		-14.82 ‰ rela	ative to VSMOW	
Tritium content of water -		na		
$\delta^{13}C$ of DIC .		na		
¹⁴ C content of DIC		na		
$\delta^{15}N$ of nitrate .		na		
$\delta^{18}O$ of nitrate .		na		
$\delta^{34}S$ of sulfate .		na		
$\delta^{18}O$ of sulfate .		na		



Water Analysis

Lab Number:	258233			Job Number:	18823
Submitter Sample Name	: Tubman-CF	5			
Submitter Sample ID:					
Submitter Job #:					
Company:	MWH Labo	ratories			
Field or Site:	OLGEP TA	SK: 401.	1.9.2		
Location:					
Depth/Formation:					
Container Type:	Amber Bott	le			
Sample Collected:	7/24/2012		Results F	Reported:	8/31/2012
δD of water		-128.8 %	‰ relative	to VSMOW	
$\delta^{18}O$ of water		-15.66 %	‰ relative	to VSMOW	
Tritium content of water		na			
$\delta^{13}C$ of DIC		na			
¹⁴ C content of DIC		na			
$\delta^{\rm 15}N$ of nitrate		na			
$\delta^{\rm 18}O$ of nitrate		na			
$\delta^{34}S$ of sulfate		na			
$\delta^{18}O$ of sulfate		na			

APPENDIX B-3

Laboratory Analytical Data - Noble Gas Data from the University of Utah

Lee Davisson

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Project Name:						Labor	tory Name
0100EF 1A3N: 401.1.9.7		T	IN MAINE	C		University of Utar	Noble Cas Laboratory
Project Contact:		3	MARAN I			Lab	Contact
Name: Victor Harris, MWH Americas, Inc.		,			Z	ame: Alan R	igby
Phone: 626-568-6024						10ne: 801-58	\$5-5214
Email: Victor.E.Harris@us.mwhglobal.com					E	nail: alan.rig	by@utah.edu
CAMPI FT D	NUMBER OF	SAMPLE LO	DCATION	SAMPLE COLLECTION	SAMPLE	ANAL	VSIS TYPE
	CONTAINERS	Northing	Easting	DATE	TIME	(Nob	le Gases)
DWP-9 T 896	2			7/23/2012		×	
DWP-9 T 898	2			7/23/2012		x	
DWP-3 T 899 V	2			7/23/2012		×	
DWP -3 T 901	2			7/23/2012		×	
DWP-6 T 911 V	2			7/23/2012		×	
DWP-7 T910 V	2			7/23/2012		x	
	Relinquished by:	M. Lee	Davissor		ă	ate/Time: 7/2	5/12 6om
	Received by:	ALD. R	2-11046	/	Ă	ate/Time: 07,	21/221
	Relinquished by:				Ă	ate/Time:	
	Received by:				ă	ate/Time:	
	Relinquished by:				D	ate/Time:	
	Received by:				Ď	ate/Time:	
	Relinquished by:				ũ	ate/Time:	
	Received by:				Ŭ	ate/Time:	

University of Utah - Dissolved and Noble Gas Lab - Sample Report

Notes	23 Large amount of excess He; poor gas model fit	Large amount of excess rie, poor gas mouer in, excess 90 air 1 amo amount of excess Har nonr nas modal fit, samul	zarge amount or excess ric, poor gas more in, sample 37 stripped	07 Large amount of excess He; OK gas model fit	Large amount or excess ris, poor gas more in, sample 46 stripped	30 Large amount of excess He; poor gas model fit	
∆Ne(%	5	1726	б <u>-</u>	36	-19	18	
Rech. Temp °c		,		16.9			
R/Ra	0.91	1.17	1.26	0.81	1.52	0.88	
He4 (ccSTP/g)	1.04E-05	3.59E-06	7.84E-05	5.63E-07	1.69E-05	4.38E-06	
Xe total (ccSTP/g)	1.17E-08	1.33E-08	1.45E-08	1.04E-08	1.28E-08	2.06E-08	
Kr total (ccSTP/g)	4.41E-08	1.26E-07	7.00E-08	3.12E-08	2.84E-08	1.12E-07	
Ne total (ccSTP/g)	1.84E-07	3.12E-06	1.57E-07	2.29E-07	1.27E-07	2.02E-07	
Ar total (ccSTP/g)	3.17E-04	1.18E-03	3.15E-04	3.68E-04	2.91E-04	3.83E-04	
N2 total (ccSTP/g)	1.16E-02	8.60E-02	3.78E-02	1.21E-02	1.05E-02	1.10E-02	
Sample I.D.	DWP-3 / T899	DWP-3 / T901	DWP-6 / T911	DWP-7 / T910	DWP-9 / T896	DWP-9 / T898	

APPENDIX C

Use of Isotope Measurements in Groundwater

Appendix C. Use of Isotope Measurements in Groundwater

C.1 Stable Isotopes

The stable isotope measurements of oxygen-18/oxygen-16 ($^{18}O/^{16}O$) and deuterium/hydrogen (D/H; deuterium is hydrogen-2) ratios in water are used to delineate different water populations in recharged groundwater. The measured $^{18}O/^{16}O$ and D/H ratios are normalized to a recognized standard and the converted results are reported in δ notation (pronounced "del"), where

$$\delta D = \left(\frac{D/H}{D/H_{std}} - 1\right) 1000$$

$$\delta^{18}O = \left(\frac{{}^{18}O/{}^{16}O}{{}^{18}O/{}^{16}O}_{std} - 1\right)1000$$

The ${}^{18}O/{}^{16}O_{std}$ and D/H_{std} are the isotopic ratios of "Standard Mean Ocean Water" (SMOW). A δ value is a per mil (or parts per thousand) deviation from the standard.

The atomic masses differences of these different isotopes in water molecules underlie differences in measured ratios. These differences arise from phase transitions in water (i.e., vapor, water, ice) which favors higher atomic masses in lower energy states of matter. For example, the measured difference in the δ^{18} O value measured between a water vapor and its condensed liquid form at 25°C is approximately 9.3 per mil. This difference is large compared to the typical measurement precision of 0.1 per mil.

The isotopic ratios of ocean water are remarkably uniform worldwide, owing to global circulation patterns. However, since all continental precipitation originates from the ocean, isotopic partitioning occurs between water phases, and because continental storm fronts are isolated from the ocean and behave as closed systems, the isotopic ratios of measured precipitation varies systematically. This variation is almost exclusively driven by elevation difference and distance inland from the ocean. An example of δ^{18} O variations in precipitation across British Columbia are illustrated below in Figure C.1a. Figure C.1b shows how shallow groundwater collected on the western slope of the Sierra Nevada record this systematic δ^{18} O variation in its recharge.



Figure C.1a,b. Figure C-1a shows systematic variation of δ^{18} O values in precipitation across British Columbia (from Yonge et al., 1989). Figure C.1b shows how shallow groundwater records this systematic variation on the western slope of the Sierra Nevada (from Rose et al., 1996).

The method for comparing the isotopic character of different waters lies in the use of a δD - $\delta^{18}O$ plot of the isotope ratios. A plot of δD vs. $\delta^{18}O$ values provides a graphical means to distinguish various populations of data relating to different water masses of different origins (Fig. C-2).



Figure C-2. General $\delta D - \delta^{18}O$ plot showing the Meteoric Water Line (MWL) and the effects of evaporation on natural waters. The slope of the evaporation line can vary between 2 and 6 and depends on the ambient temperature and humidity. The MWL has a constant slope of 8 for global precipitation.

Also on this plot lies what is referred to as the Global Meteoric Water Line (MWL), a linear regression through the values of various unevaporated precipitation collected world-wide, which results in an empirical equation of $\delta D = 8 \delta^{18}O + 10$. The slope of this line originates from the fact that isotopic partitioning of deuterium between water vapor and liquid is approximately 8 times greater than for ¹⁸O. Since global precipitation forms

a slope of 8 indicates that cloud water establishes isotopic equilibrium between vapor and liquid.

However, when liquid water evaporates from the surface of water body, a nonequilibrium partitioning develops between the relative deuterium and ¹⁸O abundances, causing isotopic enrichment of the remaining liquid water. On a δD - $\delta^{18}O$ plot, progressive evaporation causes a shift of the remaining liquid to the right of the MWL along a straight line (see Fig. C-2). The slope of this evaporation line depends on temperature and humidity of the surrounding air. The proximity of an evaporated isotopic value relative to the MWL is proportional to the extent of evaporation or isotopic enrichment.

C.2 Tritium-Helium-3 Age Dating

Attempts have been made in the past to date groundwater with the radioactive (unstable) hydrogen-3 isotope tritium (³H; see Mazor, 1991 and references therein). Because of its radioactive half-life of 12.43 years, it is ideally a good chronometer for young (\leq 40 years) groundwater flow. Unfortunately from a dating standpoint, ³H concentrations in precipitation have varied considerably over the past 30 years due to ³H production from surface testing of thermonuclear weapons (Fig. C-3).



Figure C-3. Changes in the ³H concentration in precipitation have varied over an order of magnitude due to fallout of thermonuclear-produced tritium from surface testing. IAEA/WMO (2001). Global Network of Isotopes in Precipitation. The GNIP Database. Accessible at: <u>http://isohis.iaea.org</u>

Tritium measurements in groundwater 20 years ago were useful from the standpoint of tracing the "bomb-pulse" ³H that had recharged into groundwater in the early 1960s and calculating the groundwater travel time based on the observed depth of the "bomb pulse". Today, however, much of the "bomb-pulse" is not well defined in groundwater due to ³H decay and groundwater dispersion. Tritium measurements alone cannot be used for dating

groundwater reliably because of the uncertainty in what the original ³H concentration was at the time of recharge, but it does serve the purpose of defining relatively young groundwater when it is observed.

In more recent years with the development of high-precision noble gas mass spectrometry, the radioactive decay product of ³H, helium-3 (³He), can be measured. The advantage to this lies in the dating equation, where

$$-17.9 \ x \ ln\left(\frac{{}^{3}H}{{}^{3}H_0}\right) = age \,,$$

³*H* is the concentration of the tritium at any given time, and ³*H*_o is the original tritium concentration at the time of recharge. Since the ³*H*_o has a large uncertainty due to the spatially and temporally variable "bomb pulse" tritium, the resulting age calculation will have large uncertainties. By simultaneously measuring the ³He produced by tritium decay (known as the tritiogenic ³He or ³*He*_{trit}) we can reconstruct the ³*H*_o by adding together the measured tritiogenic ³*He*_{trit} and the ³*H* which leads to

$$-17.9x ln\left(\frac{{}^{3}H}{{}^{3}H+{}^{3}H}_{trit}\right) = age,$$

Dissolved ³He measured in a groundwater is actually derived from several sources that include:

$${}^{3}He_{meas} = {}^{3}He_{trit} + {}^{3}He_{equil} + {}^{3}He_{excess} + {}^{3}He_{rad}$$

where ${}^{3}He_{meas}$ is the total 3 He analytically measured, ${}^{3}He_{equil}$ is the amount of 3 He dissolved in a non-turbulent surface water in equilibrium with the atmosphere and is temperature dependent, ${}^{3}He_{excess}$ is the amount of 3 He dissolved in water exceeding the equilibrium amount (a common phenomenon in groundwater due to excess dissolved air), and ${}^{3}He_{rad}$ is the amount of 3 He produced from radioactive decay of isotopes other than tritium. The latter species is very minor and totals only about 0.2% of the total 3 He. Separating these different components of the 3 He requires additional measurements of the 4 He abundance which comprise:

$${}^{4}He_{meas} = {}^{4}He_{equil} + {}^{4}He_{excess} + {}^{4}He_{rad},$$

where the subscripts are the same as those for ³He. In the case of ⁴ He_{rad} , a product of uranium-thorium decay, the abundance can be significant where older waters are involved (e.g. >1000 years old) and has been used numerous times as an independent groundwater age measurement due to its steady state accumulation.

The ${}^{3}He_{equil}$, ${}^{4}He_{equil}$, and ${}^{4}He_{rad}$ terms are either assumed or determined by other noble gas abundance measurements (see below), while the ${}^{3}He_{rad}$ term is assumed. The

two unknowns left are the excess air terms and the tritiogenic 3 He, of which we can formulate two equations to solve for them.

The ${}^{4}\text{He}_{\text{meas}}/{}^{4}\text{He}_{\text{equil}}$ ratios provide a method for determining the excess air contribution to the sample, since a ratio >1.0 is created by incorporation of more dissolved helium than in equilibrium with the atmosphere, assuming an appreciable amount of ${}^{4}\text{He}$ has not accumulated from radioactive decay (see below). This assumption can be validated with additional noble gas measurements. If radiogenic ${}^{4}\text{He}$ is a concern, though, the ${}^{3}\text{He}/{}^{4}\text{He}$ ratios can be calculated and compared to ratios expected in water at equilibrium concentrations. This comparison is important since if there is any appreciable radiogenic ${}^{4}\text{He}$, then the ${}^{3}\text{He}/{}^{4}\text{He}$ ratio relative to equilibrium will be <1.0. This is due to the accumulation of ${}^{4}\text{He}$ from uranium-thorium decay. Where there are indications of radiogenic ${}^{4}\text{He}$ we can correct for it in the age calculations.

C.3 Noble Gas Abundance

The noble gases of helium, neon, argon, krypton, and xenon naturally occur at trace abundance in the atmosphere. They also dissolve in groundwater during recharge. Their concentration in groundwater is controlled by 1) equilibrium solubility and 2) incorporation of excess air. The solubility of the noble gases in non-turbulent, freestanding water is temperature dependent, with increasing solubility with decreasing temperature. This temperature dependency is most pronounced in the argon, krypton, and xenon concentration (Fig. C-4).



Figure C-4. Solubility of noble gases in water at various temperatures can be used to calculate groundwater recharge temperatures. See Mazor (1991) for examples and further discussion.

The curves in figure C-4 provide a means to calibrate measured dissolved noble gas abundances in groundwater against its recharge temperature. During most groundwater recharge, the mean soil temperature dictates the equilibrium noble gas concentrations dissolved in recharging water, which in most regions is around 2°C greater than the mean annual air temperature.

Dissolved noble gas abundances in groundwater other than helium that exceed an equilibrium amount are due to dissolution of excess air. Incorporation of excess air into recharged groundwater is thought to occur when air in the vadose zone is trapped by a plug of recharge water and is transported to deep enough depths that it is dissolved.

Groundwater recharged through a vadose zone likely has excess dissolved air. In almost all cases the composition of the excess air is the same as the atmosphere (Heaton et al., 1981). Therefore, the amount of noble gases dissolved in groundwater above the equilibrium amount is a simple arithmetic addition of each noble gas from the atmosphere. Therefore, the amount of each dissolved noble gas relative to each other within a single sample should reflect a single equilibrium solubility temperature at the time of groundwater recharge. The amount of excess air dissolved in a groundwater can also provide qualitative information about the type of groundwater recharge. For instance, high excess air content may suggest recharge by a periodic "piston" flow under vadose zone conditions. Little excess air may suggest recharge with a limited vadose zone such as in river or lake infiltration.

The remaining noble gas effect that requires some consideration is the build-up of radiogenic ⁴He. There is a constant flux toward the ground surface of ⁴He derived from radioactive decay of uranium and thorium in the Earth's crust that, given enough time, can accumulate in groundwater. Typically groundwater that is thousands of years old will have an appreciable amount of radiogenic ⁴He, while young groundwater (<100 years old) has little or none except in special conditions such as close proximity to large-scale active faults.

To test for the presence of radiogenic ⁴He, the other noble gas abundances must be measured and calibrated to a recharge temperature. With this recharge temperature, the ⁴He content can be predicted based on equilibrium solubility. Any ⁴He that is above this predicted amount can be attributed to radiogenic ⁴He, and subsequently subtracted. This will provide a revised ³He/⁴He ratio that can be used for calculating the groundwater age.

C.5 Radiocarbon and Carbon-13

Radiocarbon, or carbon-14 (¹⁴C) is a radioactive isotope of carbon with a half-life of 5730 years. For decades ¹⁴C has been used for age-dating of carbon-bearing materials (e.g. archeological artifacts) in the range of 100 to 50,000 years. Groundwater has also been dated, and most commonly by the ¹⁴C abundance in dissolved inorganic carbon (DIC). Although many successful studies have been conducted using DIC ¹⁴C measurements, much debate still continues about how and to what the extent carbonate minerals in aquifer systems dilute ¹⁴C in recharging groundwater. As a result, absolute age determinations of groundwater using ¹⁴C are limited to special cases where the absence of carbonate can be demonstrated or ¹⁴C correction models can be validated. For the most part, absolute ages ≤ 1000 years old are usually highly uncertain.

The stable isotope of carbon, carbon-13 (¹³C), is often measured in DIC and can provide either a source indicator or a relative measure of carbonate mineral reaction. Groundwater acquires DIC during recharge through plant root zones. The partial pressure of CO₂ in the soil root zone is usually higher (i.e. factor of 2 to 1000) than the atmosphere. Recharging groundwater will dissolve this soil zone CO₂, which is chemically neutralized by dissolution of minerals. Soil carbonate is the most common mineral interaction, but in its absence, aluminosilicates can also serve as a reactive substrate. Atmospheric CO₂ has a δ^{13} C value of approximately –7.5 per mil (the del system is the same as used for ¹⁸O and deuterium, but carbon isotope ratios are compared to a reference carbonate material instead). Higher plants growing on the surface use this CO₂ for photosynthesis and in the process preferentially use ¹²C over ¹³C. As a result, plant δ^{13} C values tend to either be around –28 per mil, or for many grasses around –13 per mil. These same δ^{13} C values will occur in the soil zone CO₂ which originates from plant roots. Consequently, the δ^{13} C of DIC in recharging groundwater will be a mixture of the root zone CO₂ and any carbonate mineral it reacts with. To complicate matters further, for root zones where the partial pressure of CO₂ can be 10 times greater than the atmosphere, and recharging groundwater is relatively slow, isotopic exchange can occur between the DIC and the atmospheric CO₂, causing an enrichment in the δ^{13} C DIC value (partitioning between DIC and CO2 is approximately 8-10 per mil, depending on temperature). This latter complication is common to desert environments. With all these variables in the recharging groundwater, predicting the final DIC ¹⁴C and δ^{13} C values of groundwater reaching the saturated zone creates many uncertainties. As a result, it is more common to take an empirical approach and compare populations of δ^{13} C values of groundwater DIC collected in the same general vicinity, and estimate the amount of carbonate interaction and the recharge dynamics.

References

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- Mazor, E., 1991, Applied Chemical and Isotopic Groundwater Hydrology. Halsted Press: New York, 274 pgs.

APPENDIX D

Elevation-Air Temperature Data for Meteorological Sites within the Owens Valley Region

Station Name	Latitude Longitude	Elevation (feet above sea level)	T°C
Independence	36°48' 118°12'	3,910	15.5
Haiwee	36°08' 117°57'	3,830	15.4
Inyokern	35°39' 117°49'	2,440	17.7
Bishop WSO Airport	37°22' 118°25'	4,150	13.3
Wildrose Ranger Stn	36°16' 117°11'	4,110	14.9
China Lake Armitage	35°41'117°41'	2,240	17.6
Trona	35°47' 117°23'	1,700	19.3
Lodgepole	36°36' 118°44'	6,750	5.4
South Lake	37°10' 118°34'	9,630	4.1
Lake Sabrina	37°13' 118°37'	9,080	5.4
Deep Spring College	37°22' 117°59'	5,230	11.4
Giant Forest	36°34' 118°46'	6,380	8.1
Huntington Lake	37°14' 119°13'	7,020	7.0
White Mt	37°35' 118°14'	12,470	-2.5